

Multiphysics couplings and instabilities II

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More is less



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Part I: Multiphysics and Plasticity

Engineering

Multiphysics - A new name to an old concept: **physics**

Momentum balance:

$$\frac{\nabla S_{ij}}{\nabla x_j} = r \frac{\nabla v_i}{\nabla t}$$

Mass balance (species?): $\frac{\nabla r^{(a)}}{\nabla t} + \frac{\nabla r^{(a)} v_k^{(a)}}{\nabla x_k} = j_a$

Energy balance:

$$rC \frac{\partial T}{\partial t} + v_k^{(m)} \frac{\partial T}{\partial x_k} = k \frac{\partial^2 T}{\partial x_k^2} + F + Dh_w r_w$$

2nd Law ThDyn:

$$\Phi = \left(\sigma_{ij} - \sigma_{ij}^Y \right) \dot{\varepsilon}_{ij}^p - Y_a \dot{\xi}_a = \chi \bar{\sigma}_{ij} \dot{\varepsilon}_{ij}^p \geq 0$$

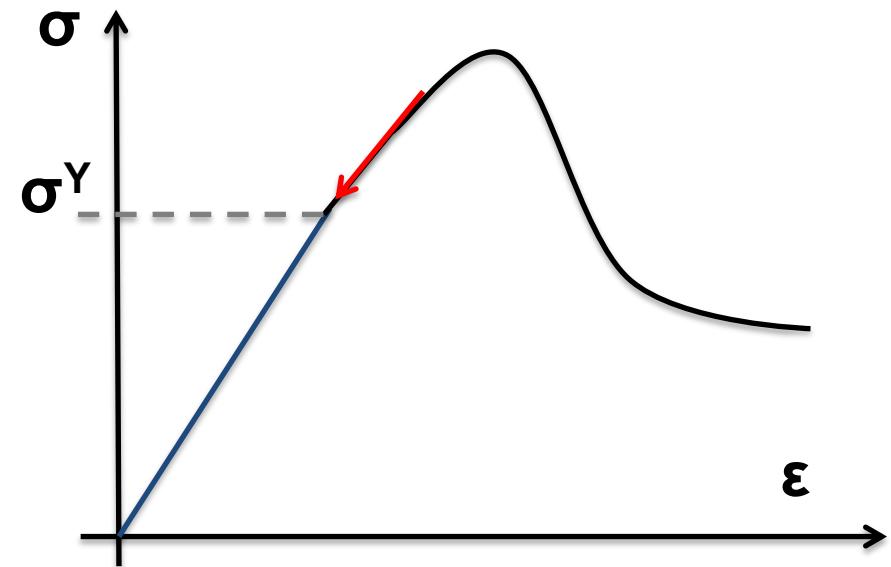


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Plasticity Theory: Consistency vs Overstress

$$\dot{\varepsilon}_{ij} = \dot{\varepsilon}_{ij}^e + \dot{\varepsilon}_{ij}^i$$

$$\dot{\varepsilon}_{ij}^i = f\left(\frac{\sigma_{ij}}{\sigma_{ref}}\right)$$

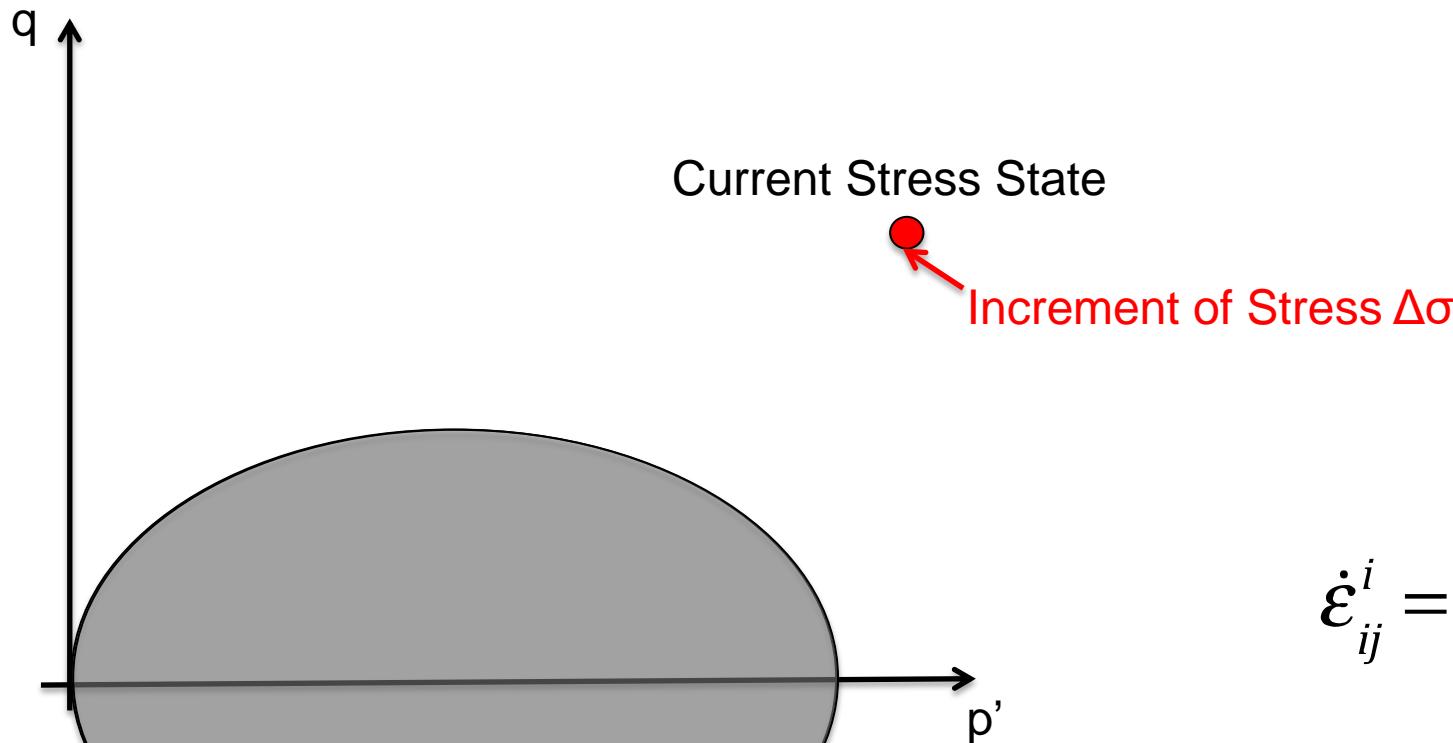


Even if the function f is not known, if it is sufficiently smooth we can expand it around the right limit of the yield surface in a Taylor series:

$$\dot{\varepsilon}_{ij}^i = f\left(\frac{\sigma_{ij}^Y}{\sigma_{ref}}\right) + \frac{\partial f}{\partial \sigma_{ij}^Y} \left\langle \frac{\sigma_{ij} - \sigma_{ij}^Y}{\sigma_{ref}} \right\rangle + \sum_{m \geq 2} \frac{\partial^{(m)} f}{\partial \sigma_{ij}^{Y,m}} \left\langle \frac{\sigma_{ij} - \sigma_{ij}^Y}{\sigma_{ref}} \right\rangle^m$$

Consistency Plasticity

$$\dot{\varepsilon}_{ij}^i = \frac{\partial f}{\partial \sigma_{ij}^Y} \left\langle \frac{\sigma_{ij} - \sigma_{ij}^Y}{\sigma_{ref}} \right\rangle + \sum_{m \geq 2} \frac{\partial^{(m)} f}{\partial \sigma_{ij}^{Y,m}} \left\langle \frac{\sigma_{ij} - \sigma_{ij}^Y}{\sigma_{ref}} \right\rangle^m$$



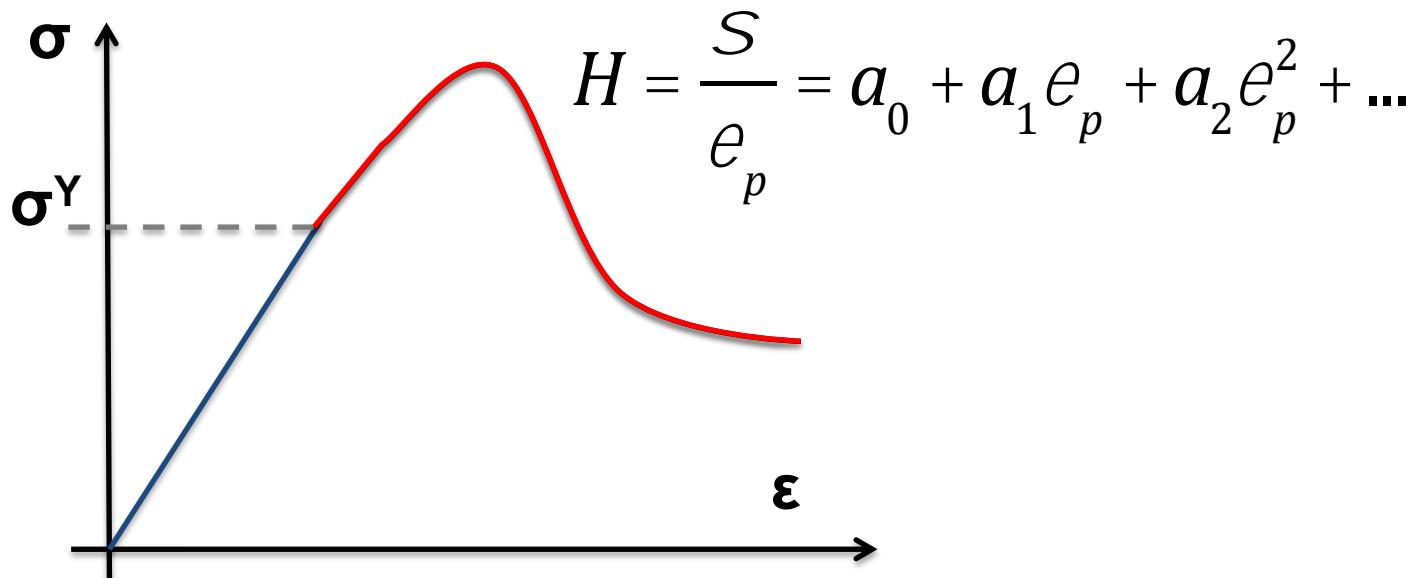
$$\dot{\varepsilon}_{ij}^i = C_{ijkl}^p \dot{\sigma}_{kl}$$



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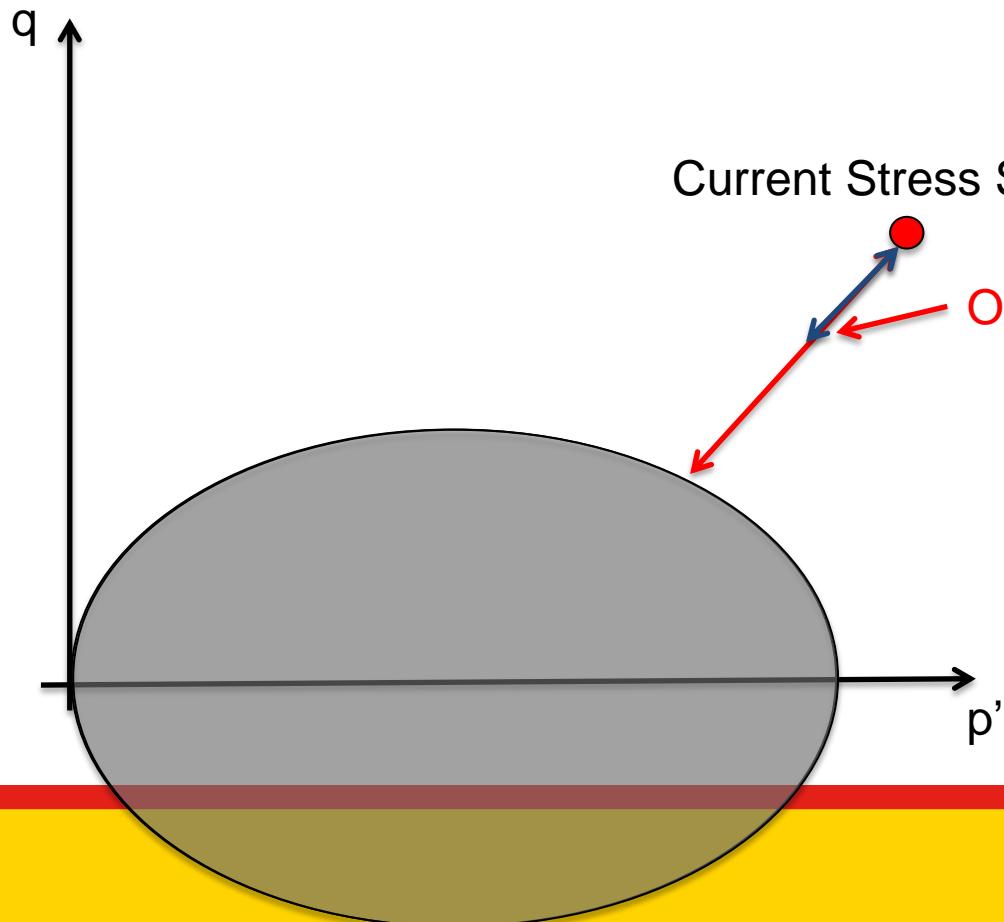
Lack of predictive power - no need in engineering design

$$\dot{\varepsilon}_{ij}^l = C_{ijkl}^p \dot{\sigma}_{kl}$$



Overstress Plasticity

$$\dot{\varepsilon}_{ij}^i = \frac{\partial f}{\partial \sigma_{ij}^Y} \left\langle \frac{\sigma_{ij} - \sigma_{ij}^Y}{\sigma_{ref}} \right\rangle + \sum_{m \geq 2} \frac{\partial^{(m)} f}{\partial \sigma_{ij}^{Y,m}} \left\langle \frac{\sigma_{ij} - \sigma_{ij}^Y}{\sigma_{ref}} \right\rangle^m$$

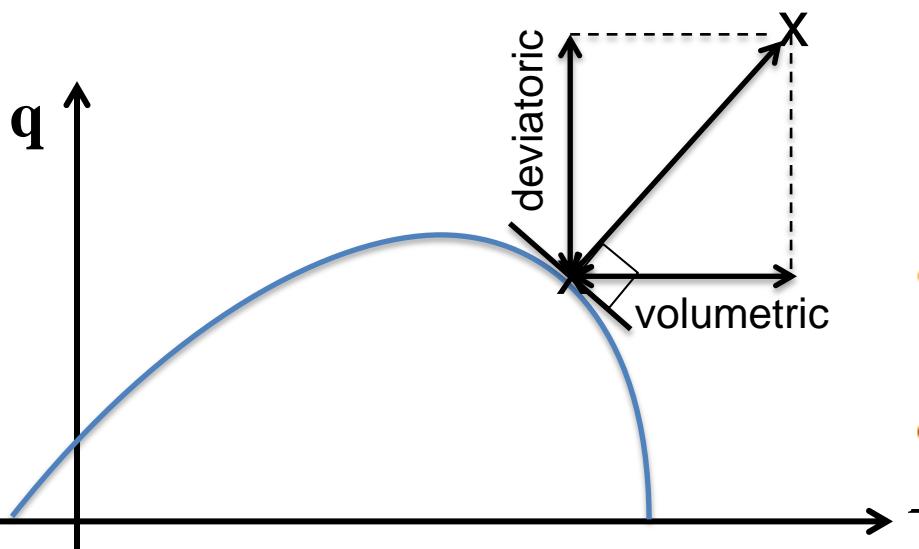


$$\dot{\varepsilon}_{ij}^i = \dot{\varepsilon}_0 \left\langle \frac{\sigma_{ij} - \sigma_{ij}^Y}{\sigma_{ref}} \right\rangle^m$$



Equations overstress plasticity: Towards a unified THMC approach

- Fluid-saturated rock
- Coaxial Elasto-visco-plasticity,
deviatoric and volumetric
components



$$\dot{\epsilon}_{ij}^i = \dot{\lambda} \frac{\partial f}{\partial \sigma'_{ij}} \quad \dot{\lambda} = \sqrt{\dot{\epsilon}_d^i{}^2 + \dot{\epsilon}_v^i{}^2}$$

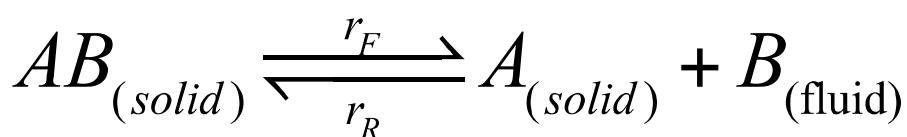
$$\dot{\epsilon}_d^i = \dot{\epsilon}_0 \left\langle \frac{q - q_Y}{\sigma_{ref}} \right\rangle^m \exp \left(-\frac{Q_{mech}^d}{RT} \right)$$

$$\dot{\epsilon}_v^i = \dot{\epsilon}_0 \left\langle \frac{p' - p_Y}{\sigma_{ref}} \right\rangle^m \exp \left(-\frac{Q_{mech}^V}{RT} \right)$$



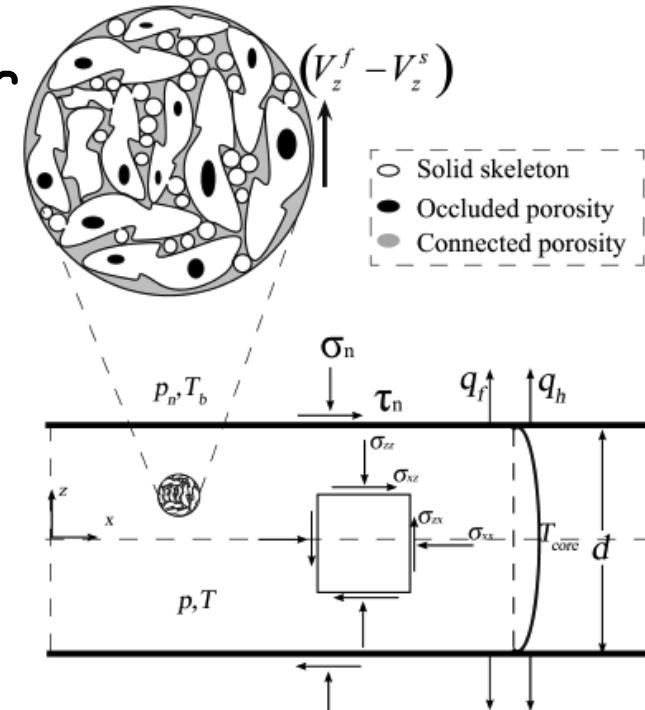
Towards a unified THMC appr

- Fluid-saturated rock
- Coaxial Elasto-visco-plasticity, deviatoric and volumetric components
- Mechanical (Shear) heating
- Endothermic fluid release reaction producing excess pore pressure



- Porosity and permeability linked with Kozeny-Carman law

$$k_\pi = k_{\pi 0} \frac{(1 - \phi_0)^2}{\phi_0^3} \frac{\phi^3}{(1 - \phi)^2}$$



$$f = f_0 + Df_{mech} + Df_{chem}$$

$$\Delta\phi_{chem} = A_\phi \frac{1 - \phi_0}{1 + \frac{\rho_B}{\rho_A} \frac{M_A}{M_B} \frac{1}{s}},$$

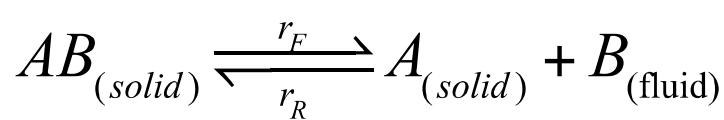
$$s = \frac{\omega_{rel}}{1 + \omega_{rel}}, \text{ and}$$

$$\omega_{rel} = \frac{\rho_{AB}}{\rho_A} \frac{M_A}{M_{AB}} K_c \exp\left(\frac{\Delta H}{RT}\right)$$

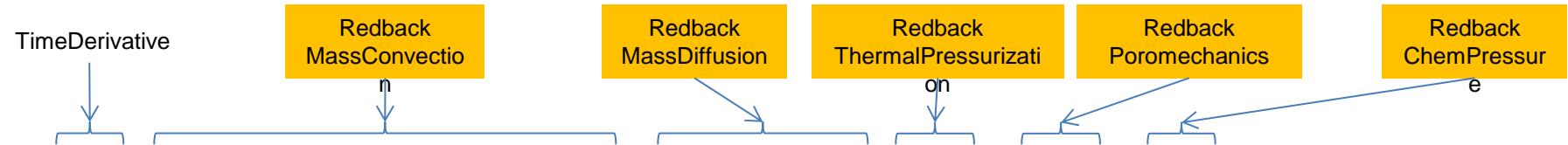


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System of equations- Redback



$$r_{F,R} = k_{F,R} r_i e^{-\frac{Ar}{1+q}}$$



$$\frac{\partial P}{\partial t} + Pe_{mass} v \cdot \nabla P - Pe_{temp} v \cdot \nabla \theta - \nabla \left[\frac{1}{Le} \nabla P \right] - \bar{\Lambda} \frac{\partial \theta}{\partial t} + \frac{\dot{\varepsilon}_v^{pl}}{\beta^*} - \mu \cdot r_F = 0$$

$$\frac{\partial \theta}{\partial t} + Pe_{thermal} v \cdot \nabla \theta - \nabla^2 \theta - X \cdot \tau \cdot \dot{\varepsilon}_d^p - X \cdot \sigma \cdot \dot{\varepsilon}_v^p + |\Delta H| \cdot r_F - |\Delta H| \cdot r_R = 0$$



Stress equilibrium

$$\frac{\partial(\sigma'_{ij} + P\delta_{ij})}{\partial x_i} = 0$$

RedbackMech Material

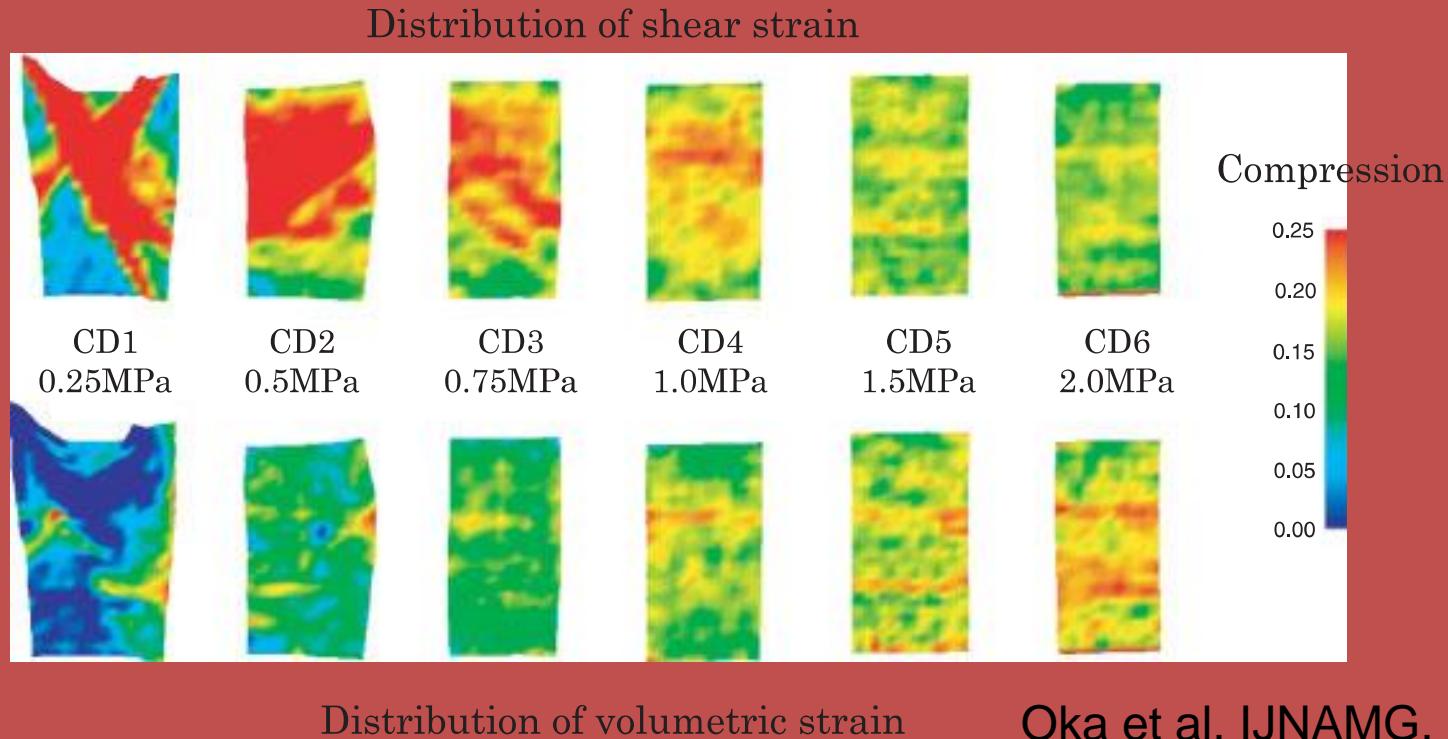
$$\dot{\varepsilon}_{ij} = C_{ijkl}^{-1} \dot{\sigma}_{kl} + \lambda \frac{\partial f}{\partial \sigma_{ij}}$$



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Calibrating the mechanics

Triaxial experiments in soft rocks



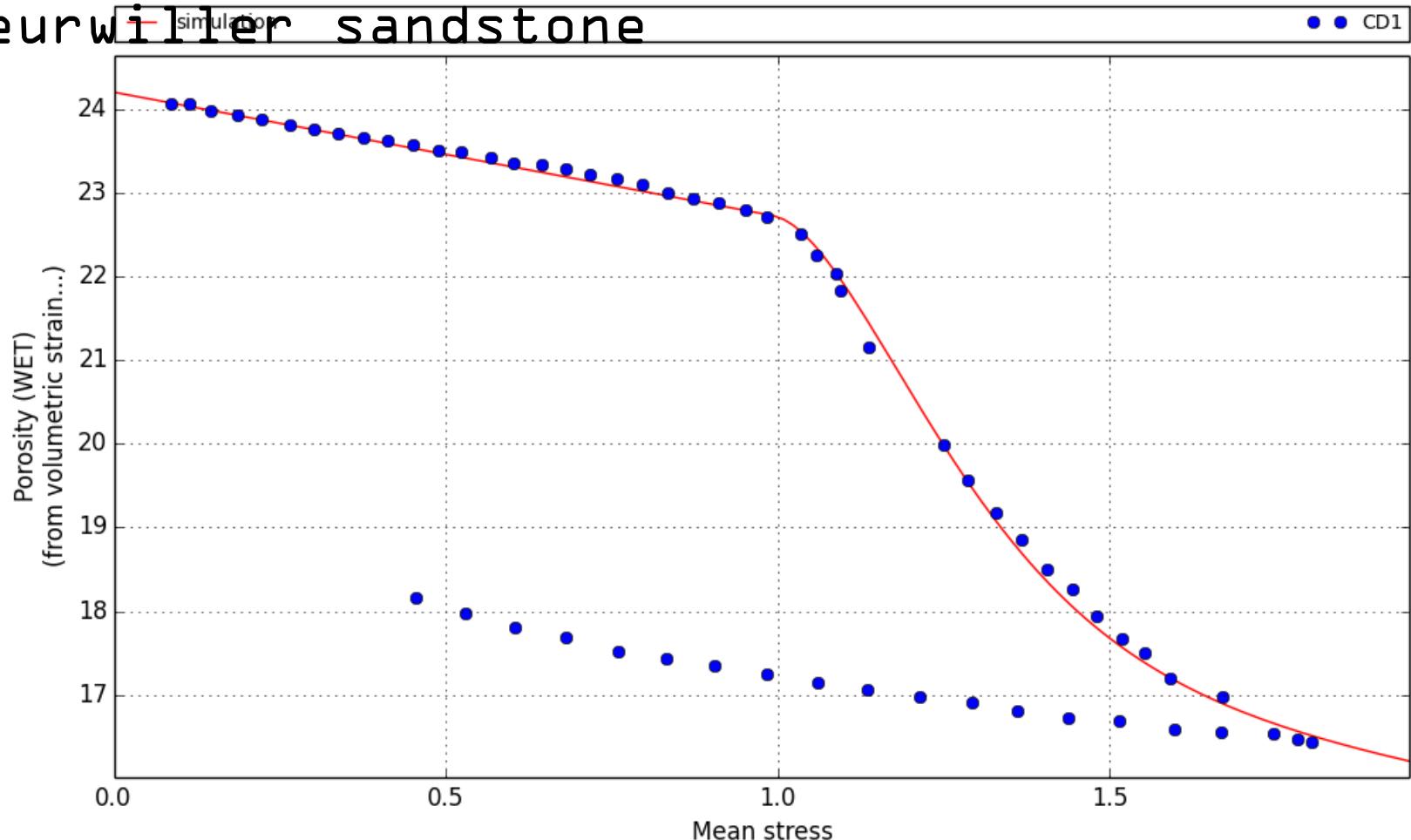
The goal is to constrain the hardening law:

$$Q_{mech} = E + p_f V_{act}$$

$$E = E_0 + DE$$

Diagenesis (pore collapse) in sandstone

Isotropic consolidation of saturated Bleurwiller sandstone



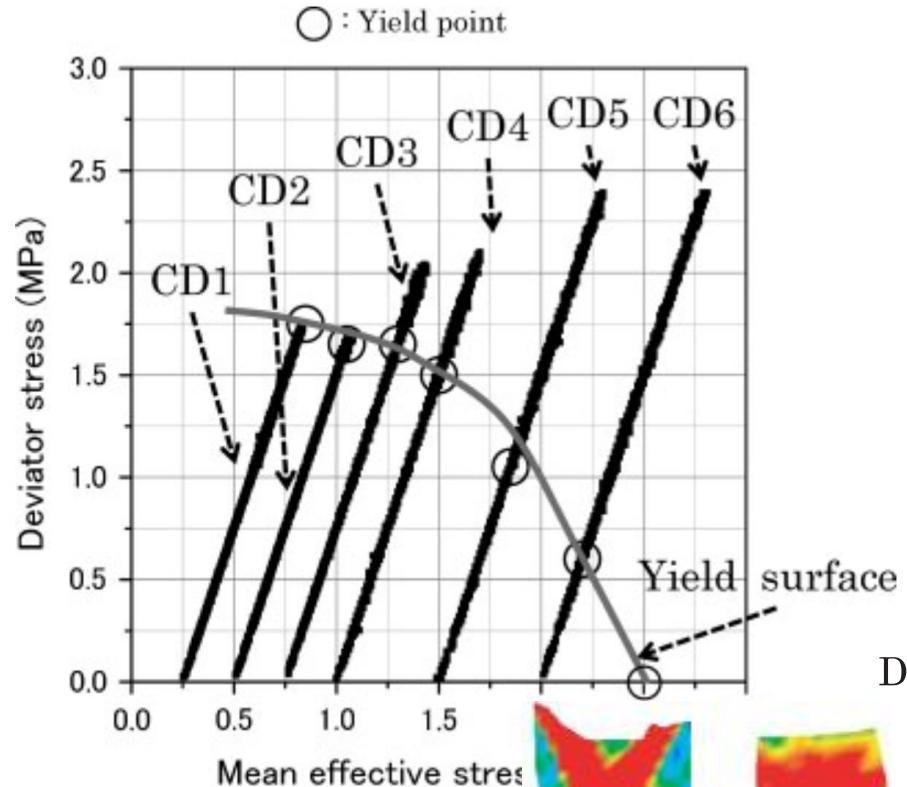
$$\Delta f_{\text{nonint}}^{\text{pores}} = p \frac{3(1 - \nu_o)}{4E_o} \left\{ \frac{10(1 + \nu_o)}{7 - 5\nu_o} \text{tr}(\sigma \cdot \sigma) - \left[\frac{1 + 5\nu_o}{7 - 5\nu_o} + \frac{1}{3(1 + \delta_s)} \right] (\text{tr } \sigma)^2 \right\},$$

Fortin et al, JGR 2007



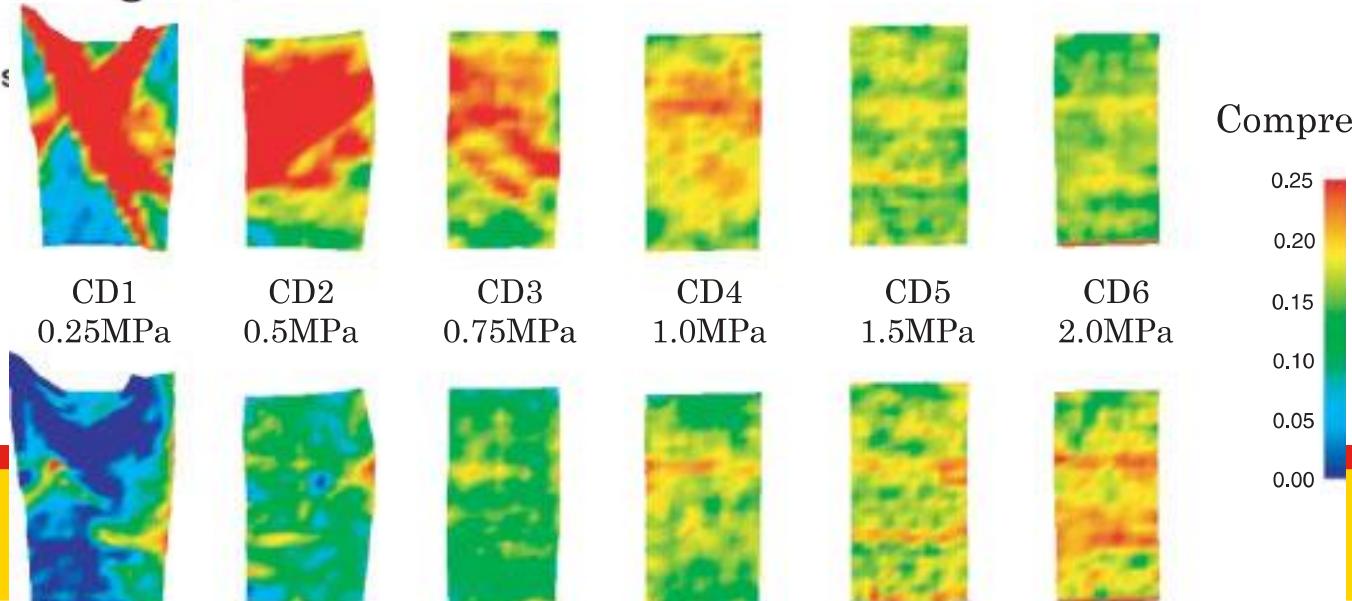
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Drained consolidation of diatomaceous

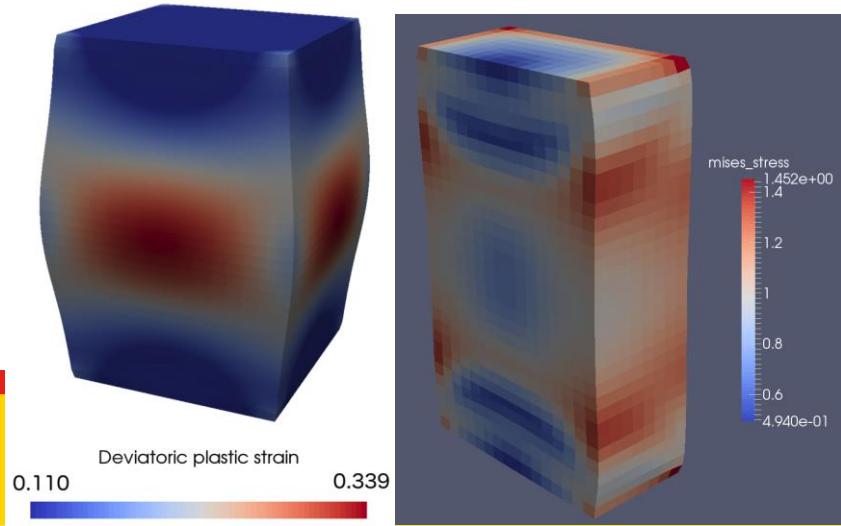
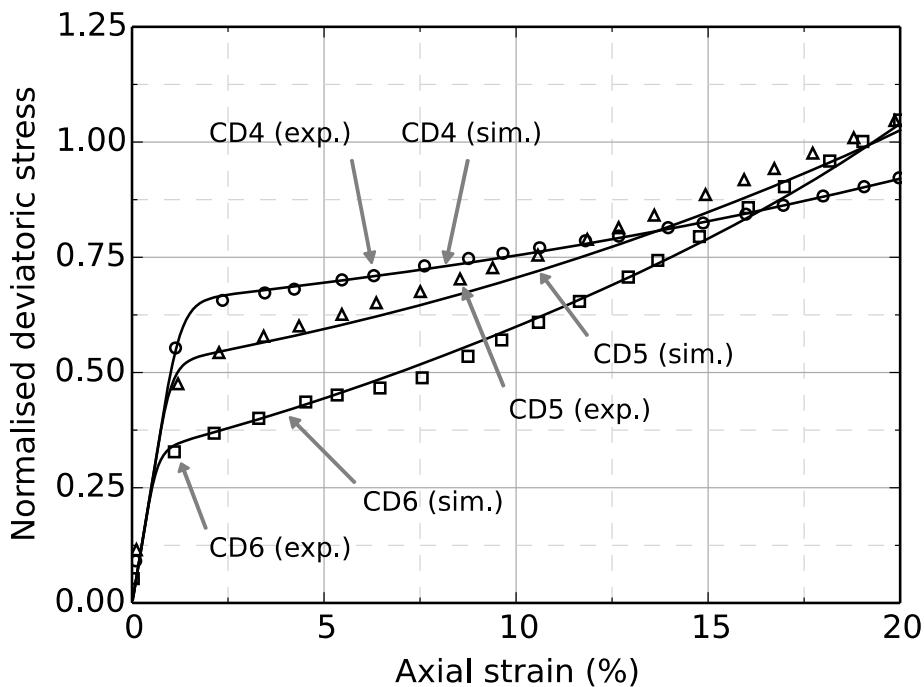
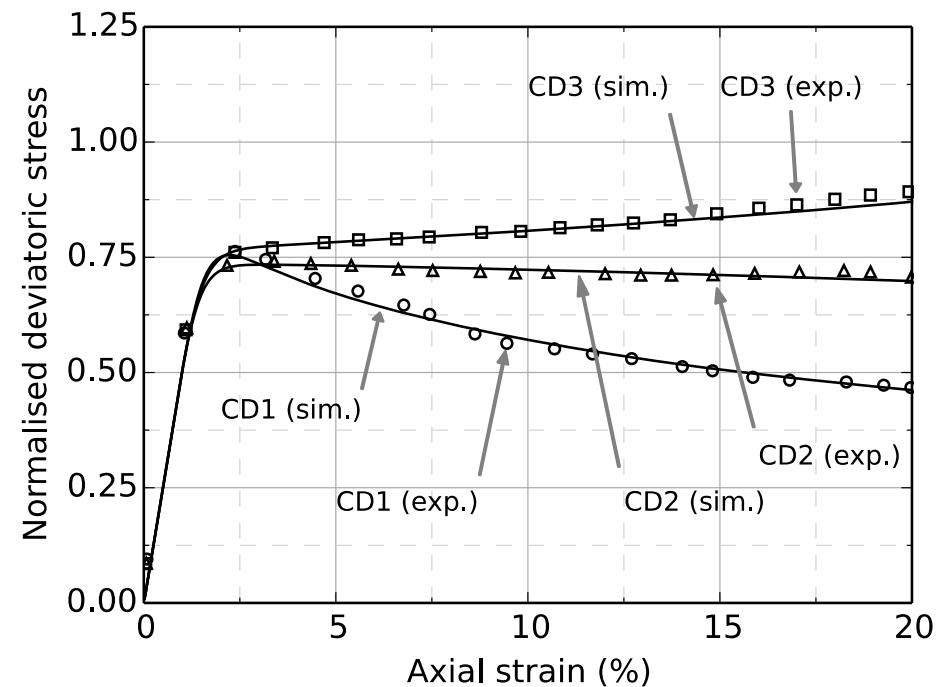


Oka et al, IJNAMG, 2011

Distribution of shear strain



Matching the experiments



$$\frac{n P_c}{n P_{cs}} \Big)$$

CD6

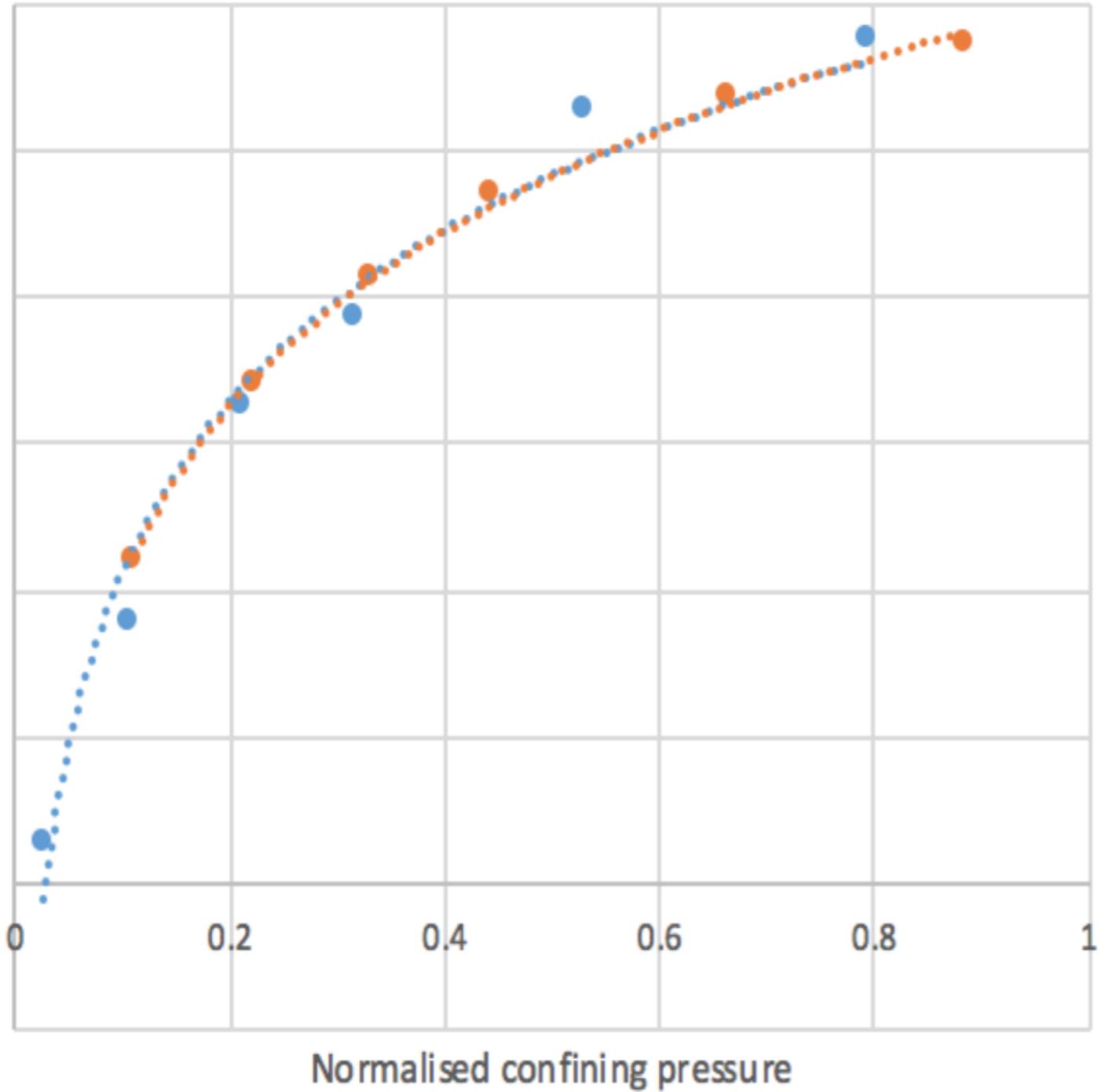
85

Sandstone

Mudstone

regime

0



Normalised confining pressure



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Part II: Bifurcation Analysis

Engineering

Equations and assumptions in multiphysics plasticity

Momentum balance:

$$\frac{\nabla S_{ij}}{\nabla x_j} = r \frac{\nabla v_i}{\nabla t}$$

Mass balance (species):

$$\frac{\nabla r^{(a)}}{\nabla x_k} = j_a$$

Energy balance:

$$\frac{\partial T}{\partial t} + v_k^{(m)} \frac{\partial T}{\partial x_k} = k \frac{\partial^2 T}{\partial x_k^2} + F + D h_w r_w$$

2nd

$$\Phi = \bar{\sigma}_{ij} \dot{\varepsilon}_{ij}^p - Y_a \dot{\xi}_a = \chi \bar{\sigma}_{ij} \dot{\varepsilon}_{ij}^p \geq 0$$

Loss of ellipticity conditions?



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Equations and assumptions in multiphysics plasticity

Momentum balance:

$$\frac{\nabla S_{ij}}{\nabla x_j} = \nabla \frac{\nabla v_i}{\nabla t}$$

Required, but
not sufficient in
these systems!!

Mass balance (species):

$$\frac{\nabla r^{(a)}}{\nabla t} + \frac{\nabla r^{(a)} v_k^{(a)}}{\nabla x_k} = j_a$$

Energy balance:

$$rC \frac{\partial T}{\partial t} + v_k^{(m)} \frac{\partial T}{\partial x_k} = k \frac{\partial^2 T}{\partial x_k^2} + F + Dh_w r_w$$

2nd Law ThDyn: $\Phi = \chi \bar{\sigma}_{ij} \dot{\varepsilon}_{ij}^p \geq 0$

Key parameter for
onset of bifurcation

Bifurcation analysis

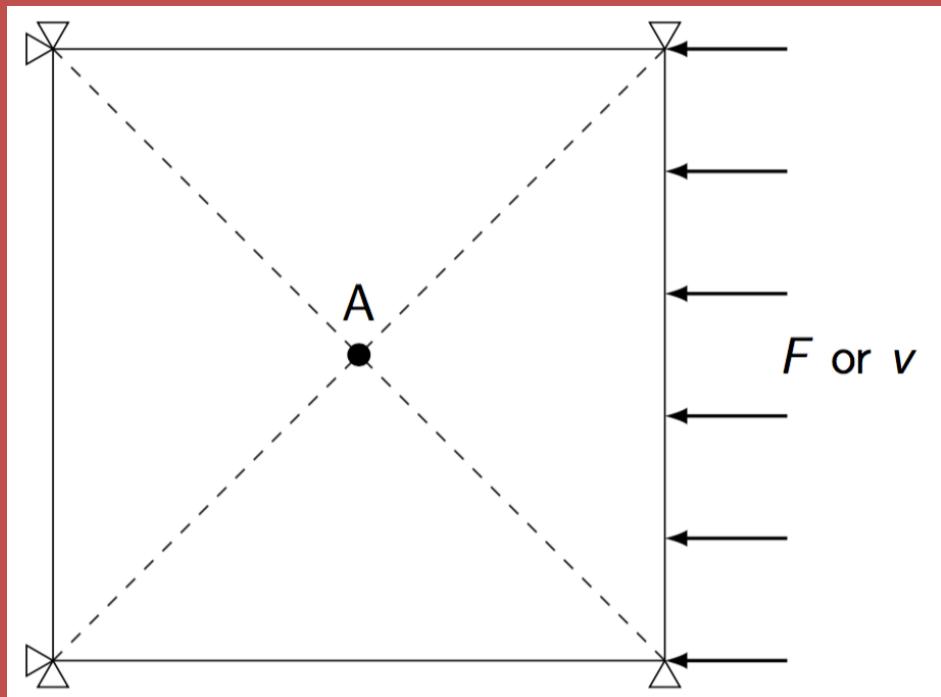
Bifurcation analysis aims on answering the following 3 questions:

1. **What** is the orientation of localized band(s)?
2. **When** does localization begin?
3. **Where** does localization begin?

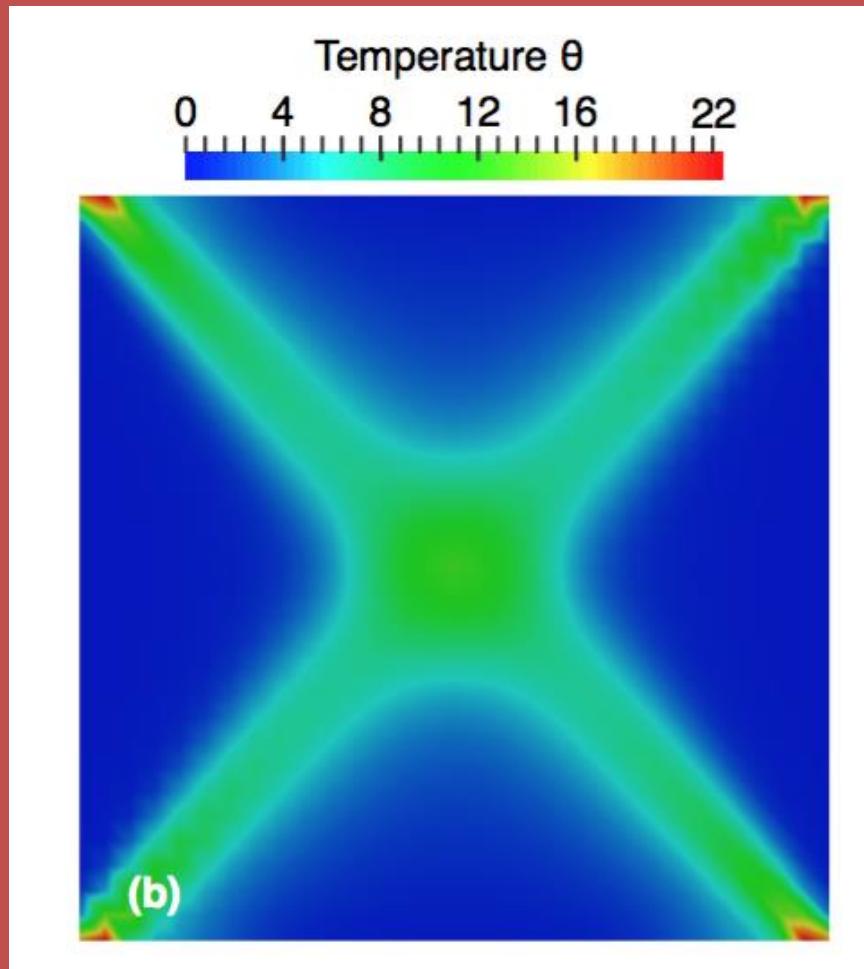


Shear bands in perlite, Milos island Greece
Photo courtesy of I. Vardoulakis

Bifurcation in Thermo-mechanical systems



Paesold et al JoMMS 2016



Thermo-mechanics: Equations

Momentum balance:

$$r \frac{\nabla v_i}{\nabla t} = \frac{\nabla S_{ij}}{\nabla x_j}$$

Energy balance:

$$rC \frac{\nabla T}{\nabla t} = k \frac{\nabla^2 T}{\nabla x_k^2} + F$$

$$\Phi = \chi \bar{\sigma}_{ij} \dot{\varepsilon}_{ij}^i = \chi \left(\bar{q} \dot{\varepsilon}_d^i + \bar{p}' \dot{\varepsilon}_v^i \right)$$

$$\dot{\epsilon}_d^i = \dot{\epsilon}_0 \left\langle \frac{q - q_Y}{\sigma_{ref}} \right\rangle^m \exp \left(-\frac{Q_{mech}^d}{RT} \right)$$

$$\dot{\epsilon}_v^i = \dot{\epsilon}_0 \left\langle \frac{p' - p_Y}{\sigma_{ref}} \right\rangle^m \exp \left(-\frac{Q_{mech}^V}{RT} \right)$$

Thermo-mechanics: Steady state analysis (fixed points)

Momentum balance:

$$\frac{\P S_{ij}}{\P x_j} = 0$$

Energy balance: $k \frac{\partial^2 T}{\partial x_k^2} + \chi \left(\langle \bar{p}' \rangle^m + \langle \bar{q} \rangle^m \right) \dot{\varepsilon}_0 e^{-\frac{Q_{mech}}{RT}} = 0$

Normalising the equations is essential!!

$$S^* = \frac{S}{S_{ref}}, \quad x^* = \frac{x}{L_{ref}}, \quad q = \frac{T - T_{ref}}{T_{ref}}$$

Thermo-mechanics: Steady state analysis (fixed points)

Momentum balance:

$$\frac{\P S_{ij}^*}{\P x_j^*} = 0$$

Energy balance:

$$\frac{\P^2 q}{\P x_k^{*2}} + Gre^{\frac{Ar q}{1+q}} = 0$$

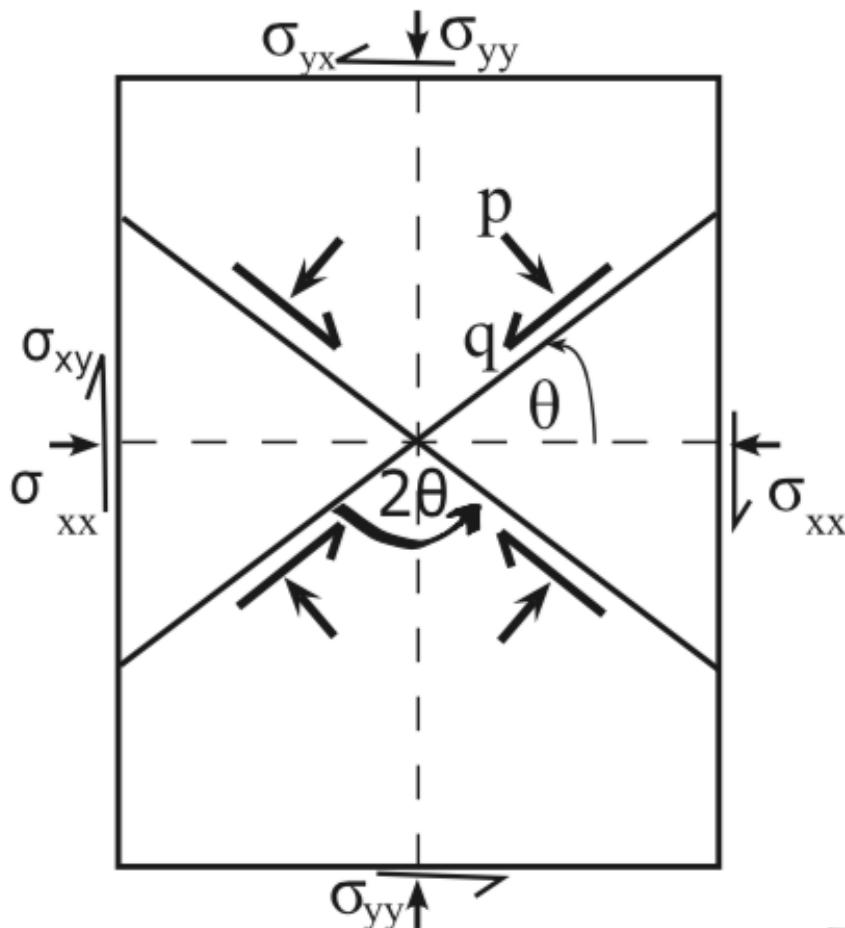
$$Gr = \frac{\Phi_0 L_{ref}^2}{k T_{ref}} \quad , \quad \Phi_0 = \chi \left(\langle \bar{p}' \rangle^m + \langle \bar{q} \rangle^m \right) \dot{\varepsilon}_0 e^{-Ar}$$

$$Ar = \frac{Q_{mech}}{R T_{ref}}$$

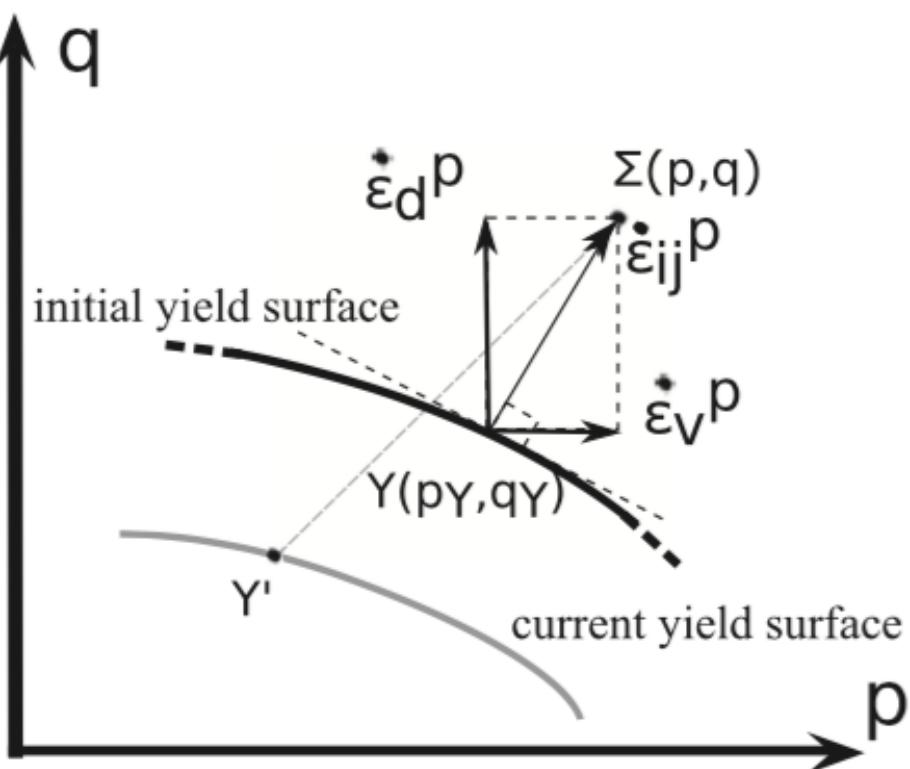
$$\chi = 1 - \frac{Y_a \dot{\xi}_a}{\bar{\sigma}_{ij} \dot{\varepsilon}_{ij}^p} ; \quad 0 \leq \chi < 1$$

Orientation of the potential localisation patterns

a



b

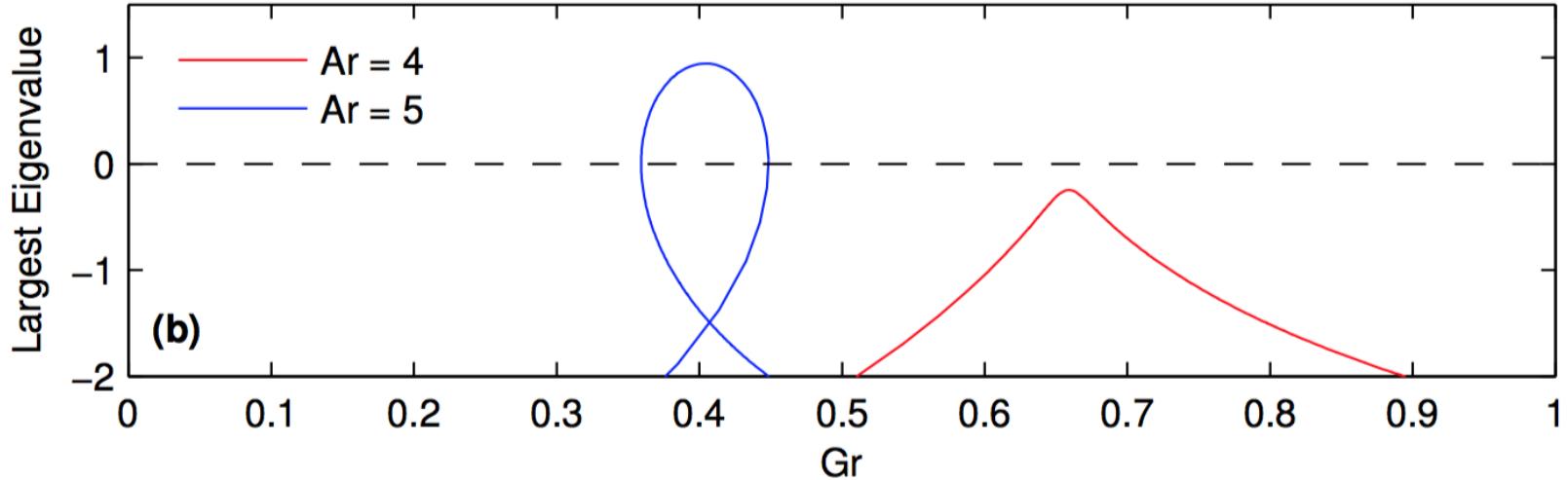
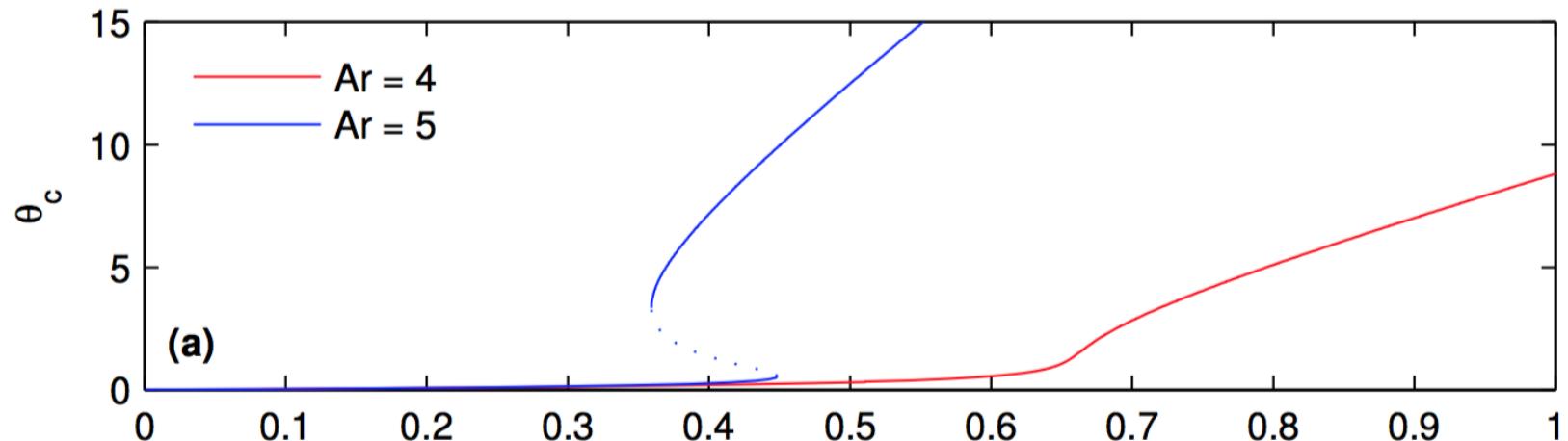


$$\frac{\|S_{ij}\|}{\|x_j\|} = 0$$



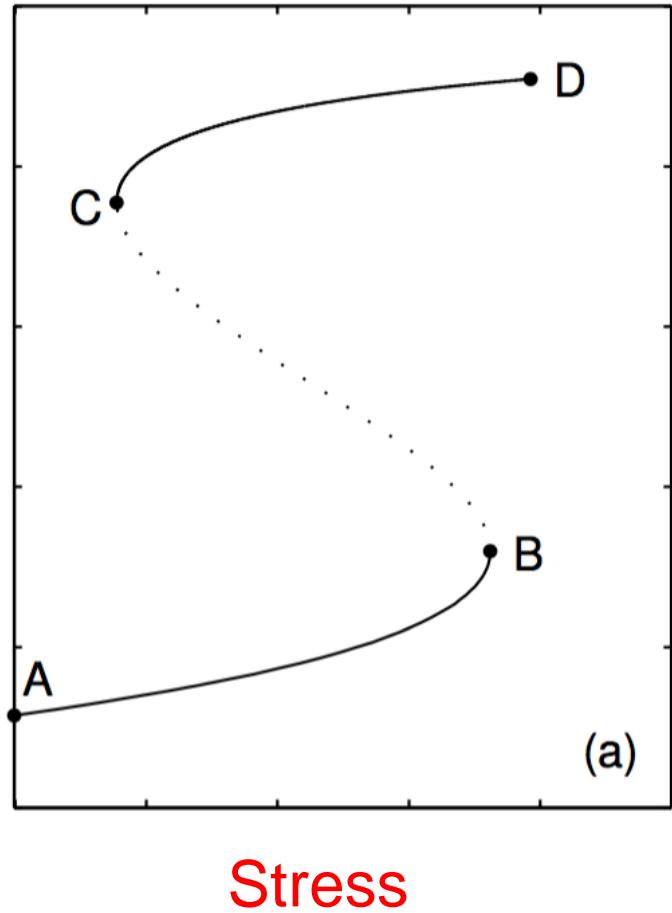
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Thermo-mechanics: Bifurcation analysis



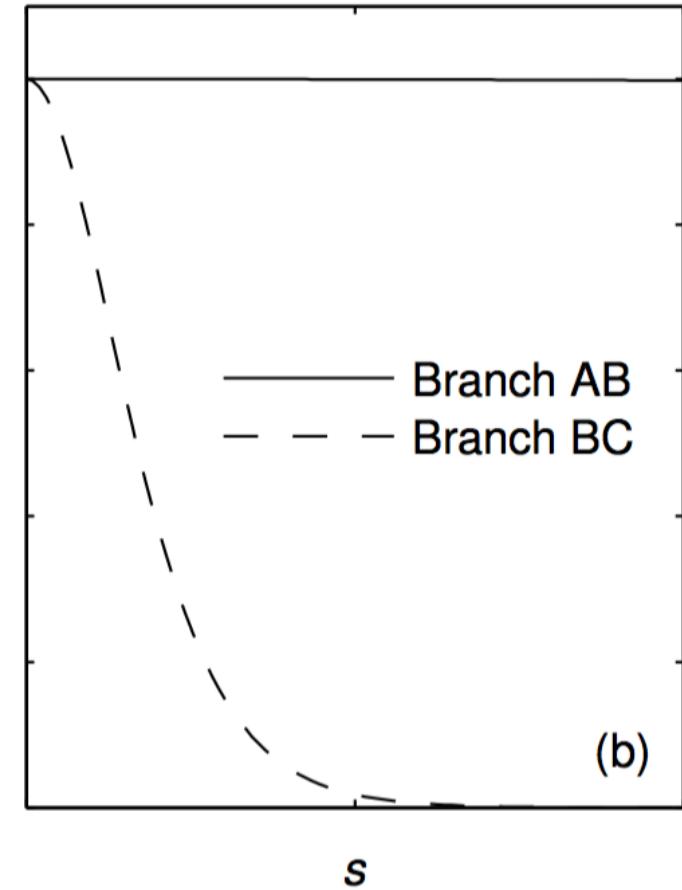
Thermo-mechanics: Bifurcation analysis

Plastic Strain Rate



Stress

ϕ/ϕ_{\max}

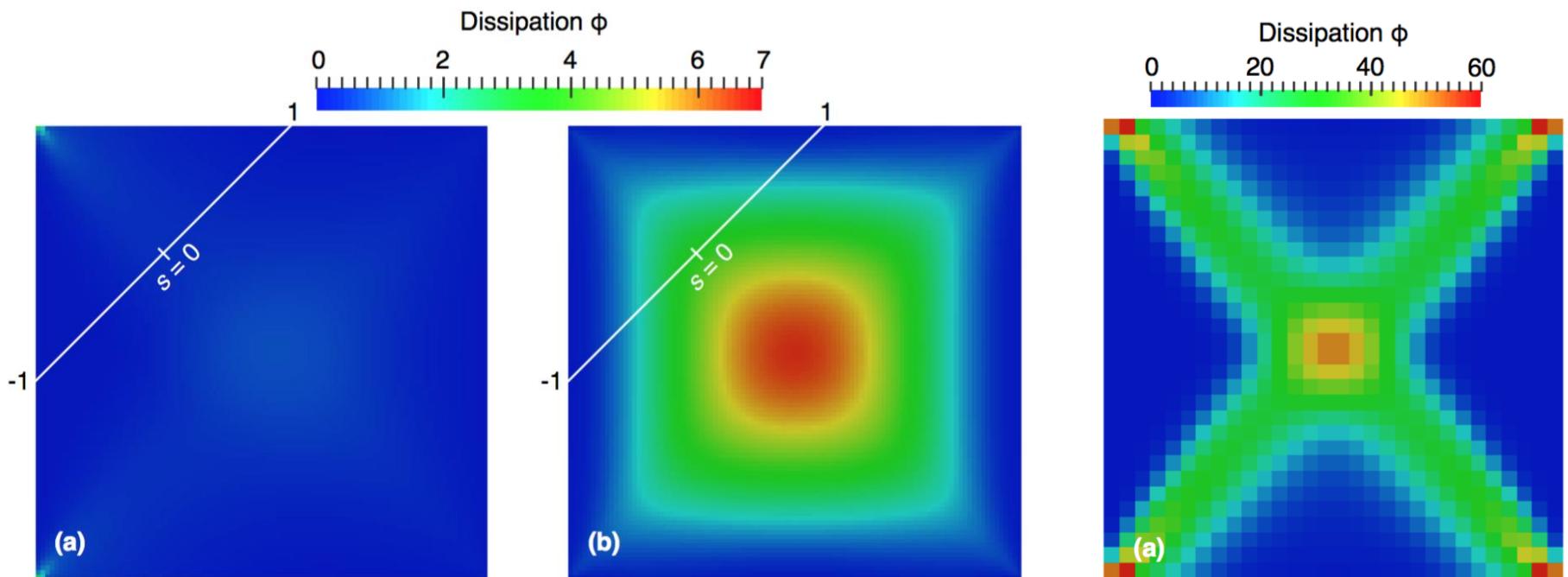


Paesold et al JoMMS 2016

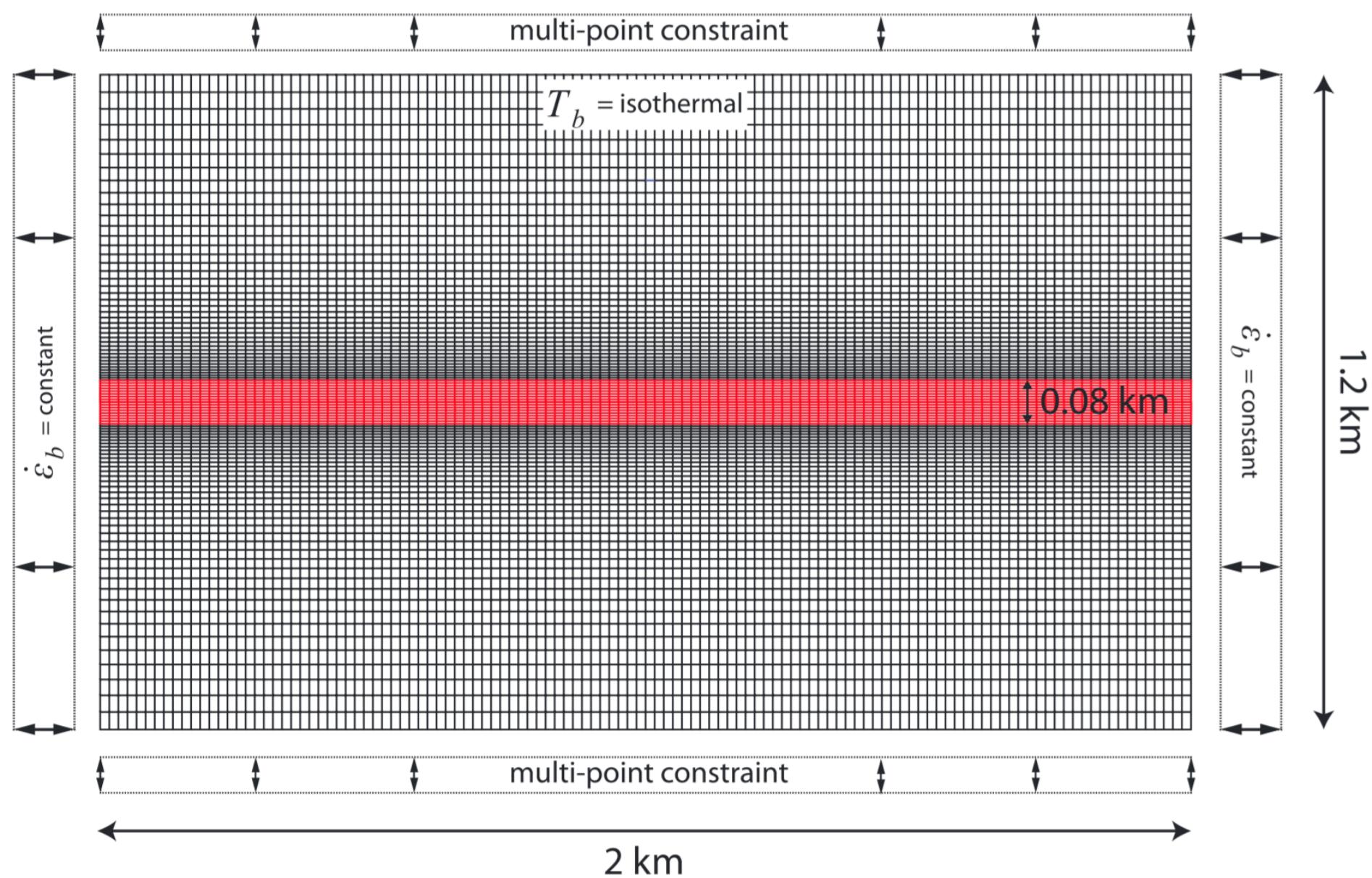


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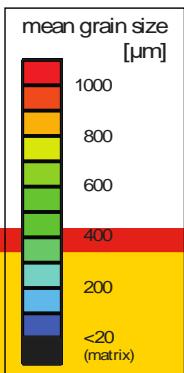
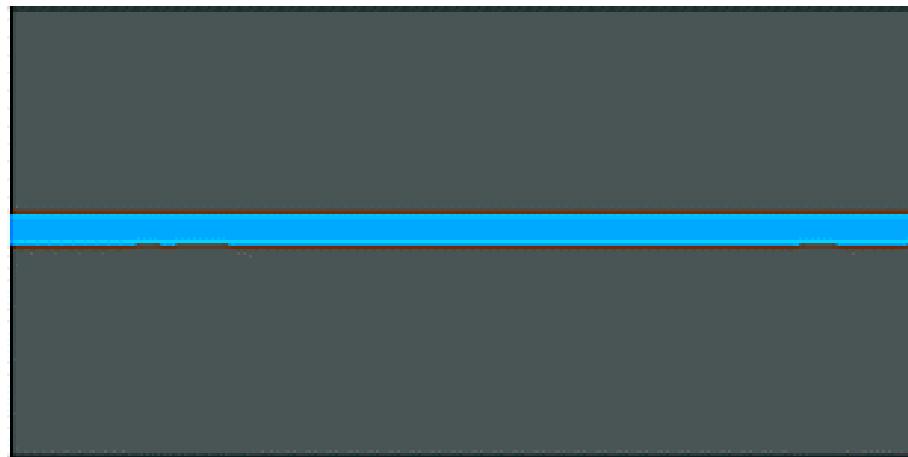
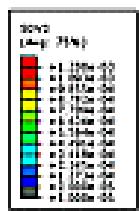
Spatial representation of localisation instability



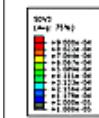
Paesold et al JoMMS 2016



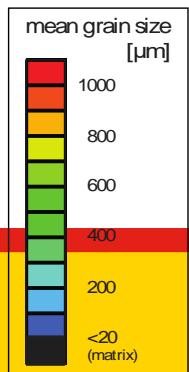
Steps: Step-1 Frame: 117
Total Time: 2003T194T1984T



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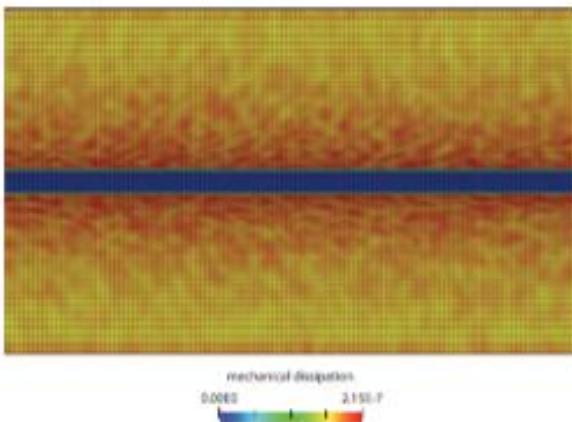
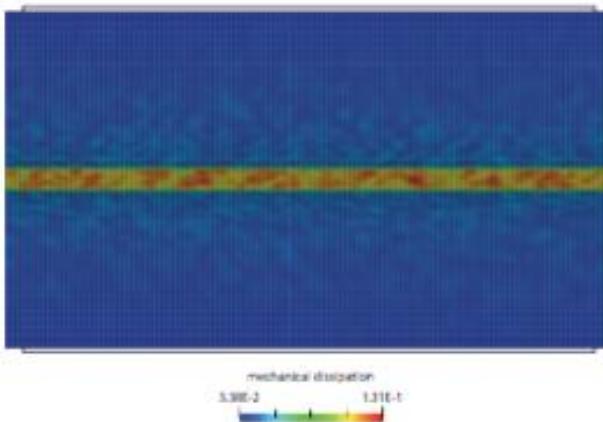
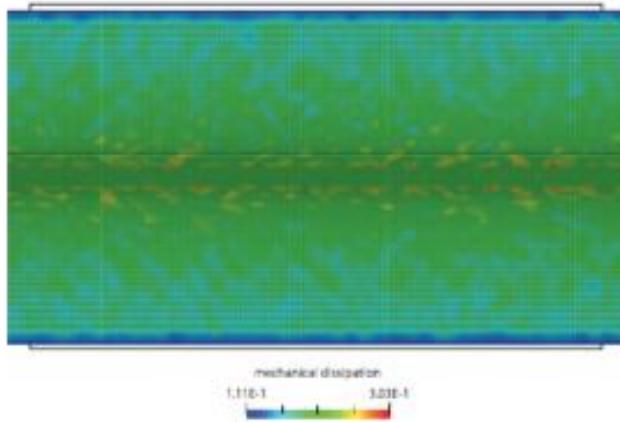


Steps: Step-1 , Frame: 105
Total Time: 2019-09-02 21:12

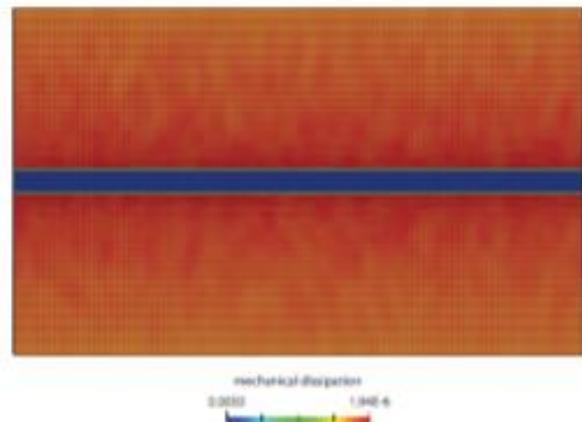
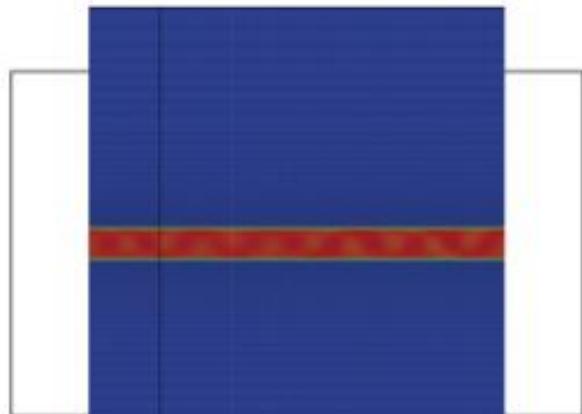


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(a)

 $\varepsilon = 1\%$ mechanical dissipation
0.0000 2.15E-7 $\varepsilon = 5\%$ mechanical dissipation
3.30E-2 1.31E-1 $\varepsilon = 7\%$ mechanical dissipation
1.11E-1 3.03E-1

(b)

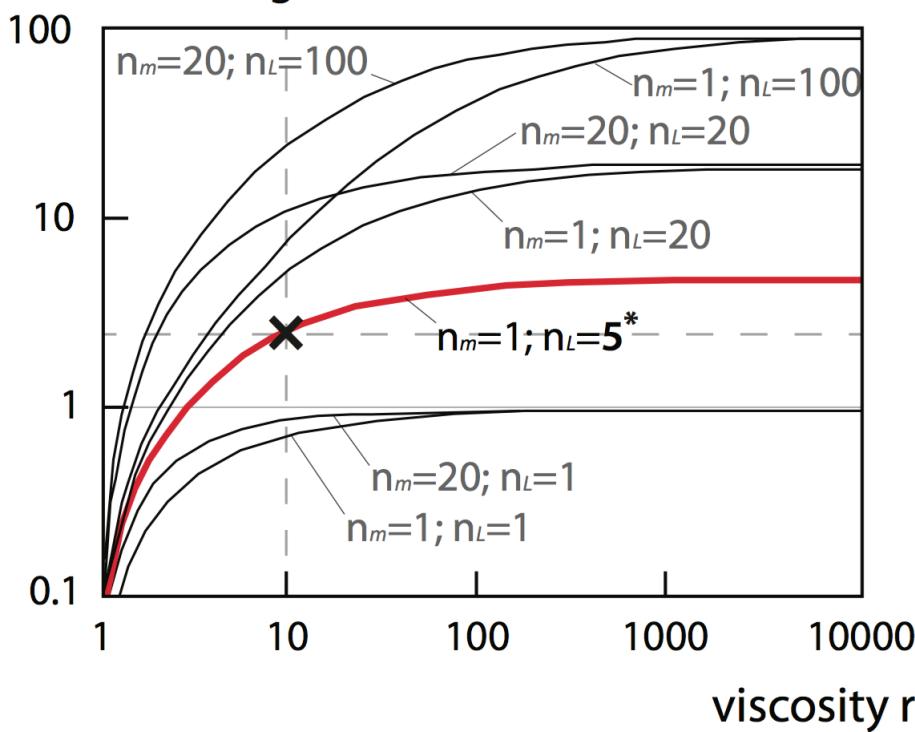
 $\varepsilon = 1\%$ mechanical dissipation
0.0000 1.04E-6 $\varepsilon = 10\%$ mechanical dissipation
4.70E-5 1.00E-2 $\varepsilon = 25\%$ 

Classical approach - different material properties required

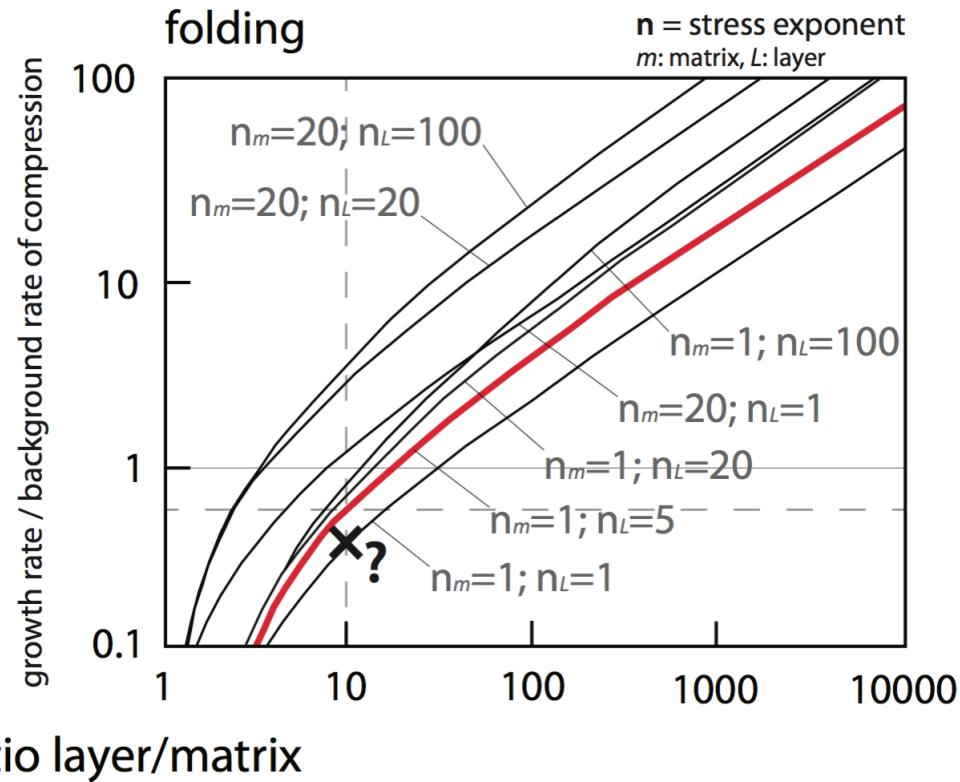
classical linear stability studies ... have some drawbacks!

e.g. Smith (1977): "solid mechanical" (plastic) materials

boudinage

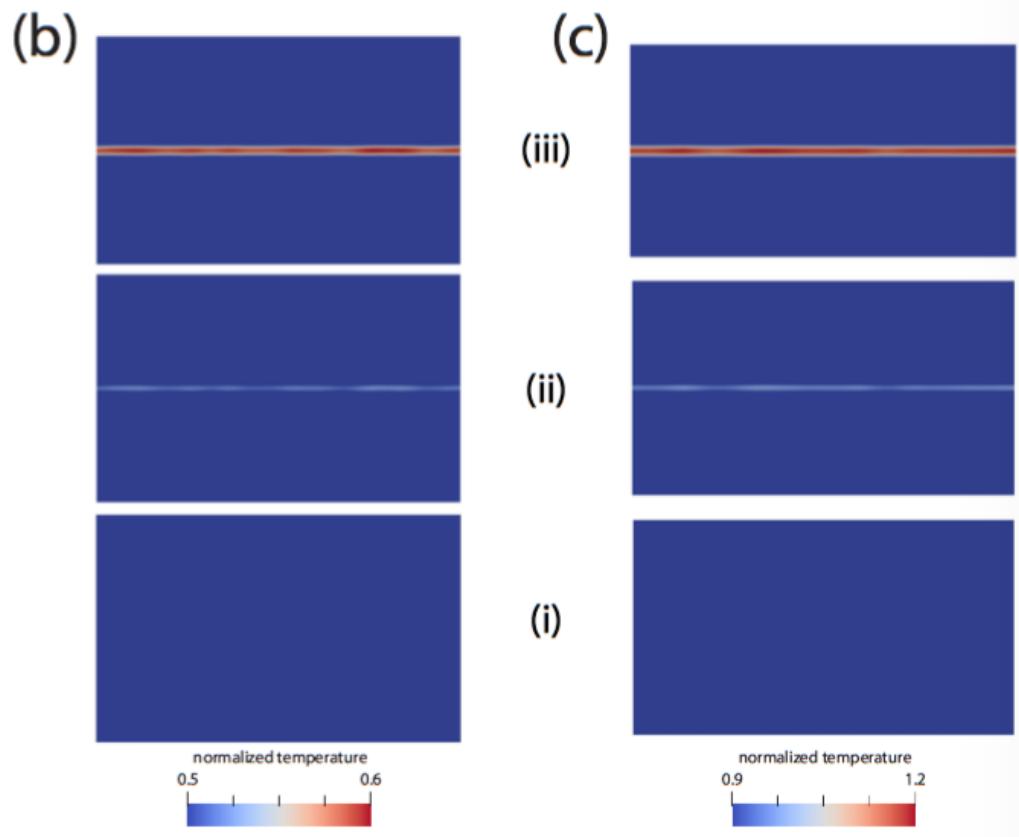
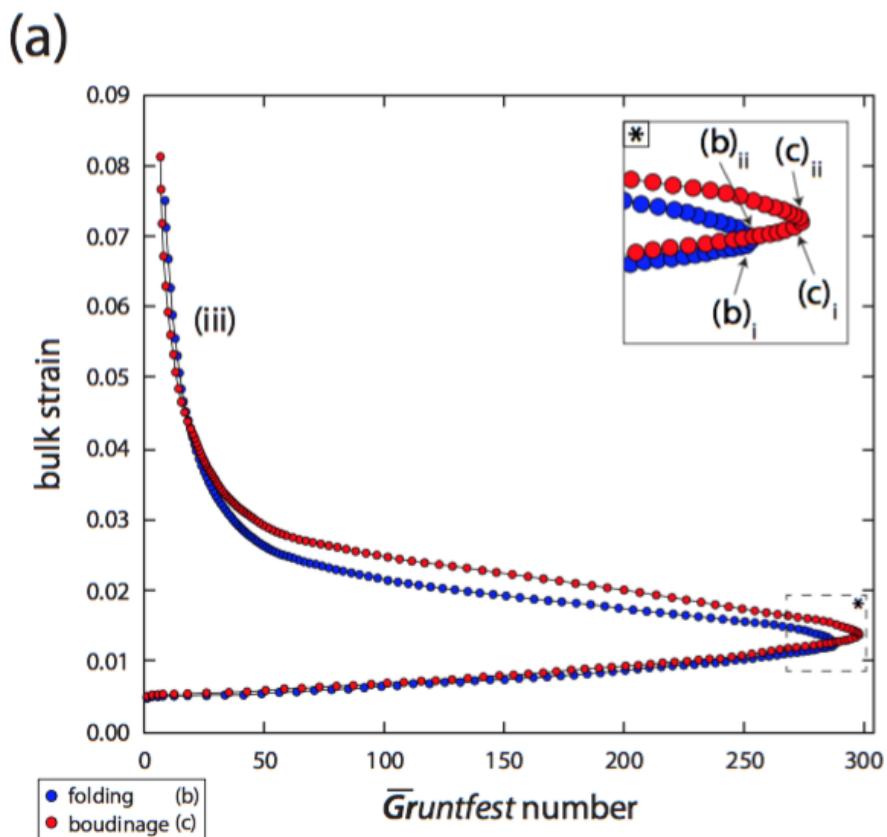


folding



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Folding and boudinage as the same energy attractor!



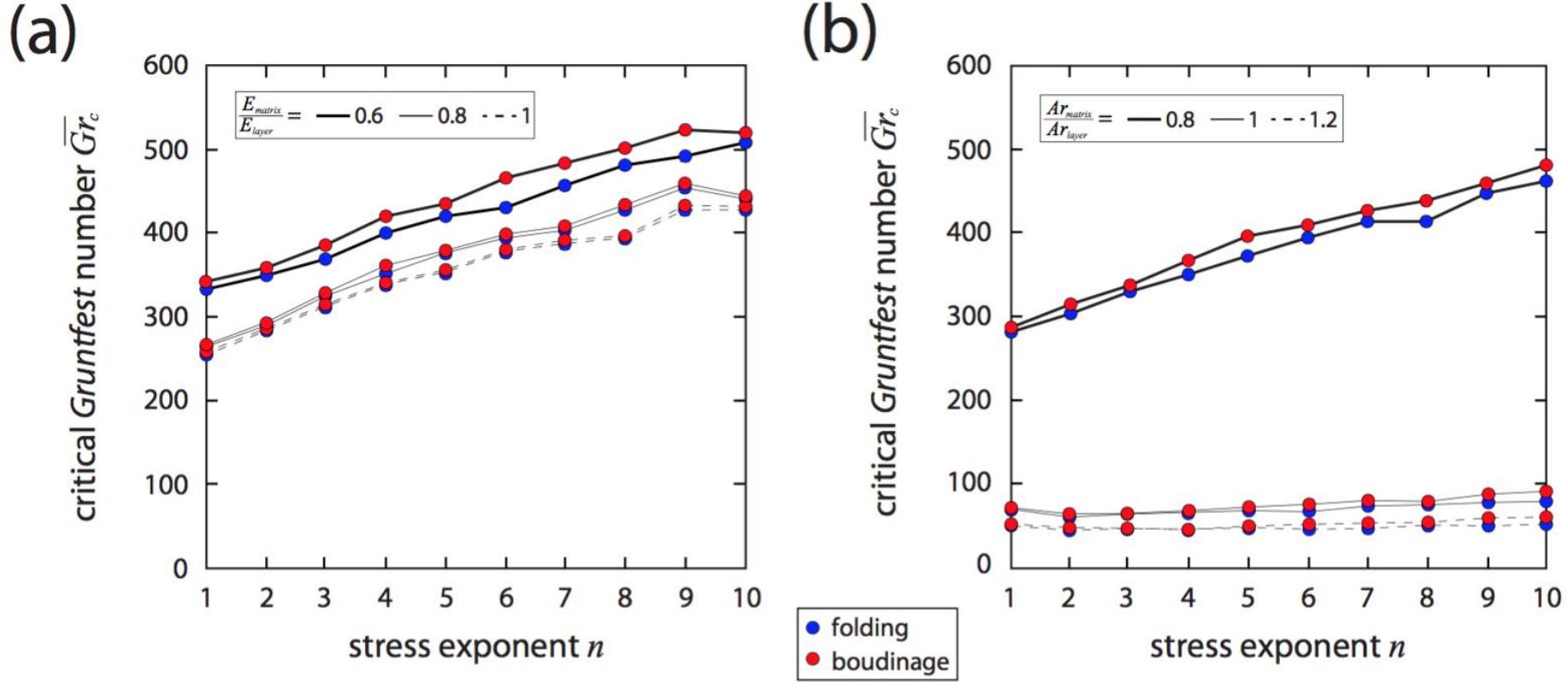
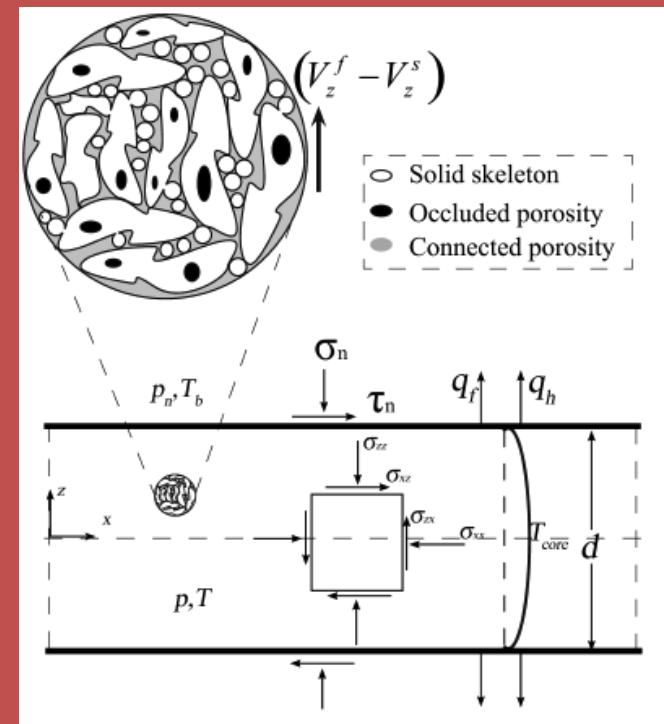


Figure 7. Parametric study of the influence of thermo-mechanical properties and the stress exponent n on the modified Gruntfest number \bar{Gr} . For all studies, the matrix obeys a linear rheology. **(a)** Evolution of \bar{Gr} for changing ratios of Young's moduli E ; constant $Ar = 8$. **(b)** Evolution of \bar{Gr} for varying ratios of the Arrhenius number Ar ; constant $E_{matrix}/E_{layer} = 0.8$. For (a) and (b) we find that \bar{Gr} follows a folded instability curve. The value of \bar{Gr}_c is almost identical for any non-linear rheology of the central layer. The energy bifurcation threshold is therefore a fixed value for any non-linear layer rheology. For thermo-mechanical parameters refer to Table 1a.

Bifurcation in THMC systems: A Homoclinic bifurcation



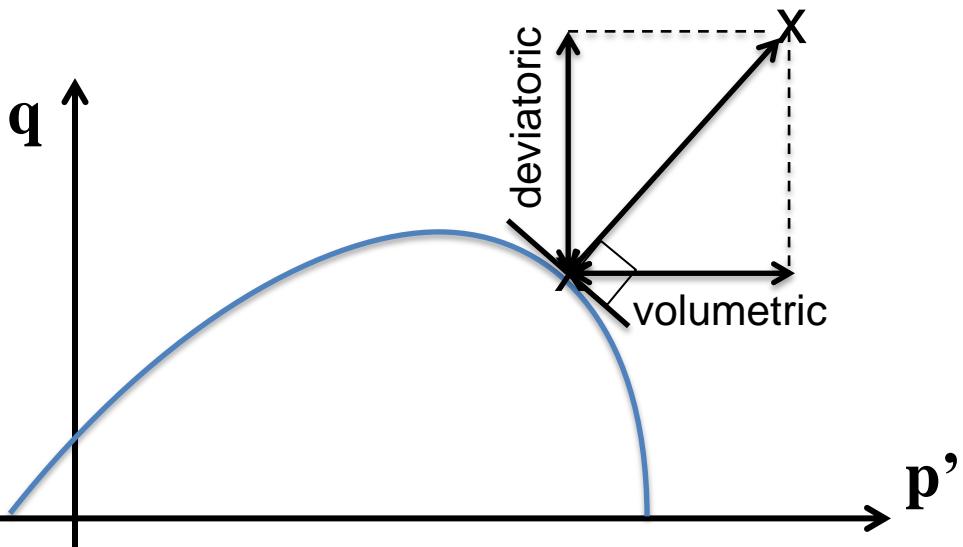
Remember the constitutive law

$$\dot{\epsilon}_{ij}^p = \dot{\lambda} \frac{\partial f}{\partial \sigma'_{ij}}$$

$$\dot{\lambda} = \sqrt{\dot{\gamma}^{p2} + \dot{\epsilon}^{p2}}$$

$$\dot{\gamma}^p = \dot{\gamma}_o \left(\frac{q - q_Y}{\sigma_{ref}} \right)^m \exp \left[-\frac{Q_{mech}}{RT} \right]$$

$$\dot{\epsilon}^p = \dot{\epsilon}_o \left(\frac{p' - p'_Y}{\sigma_{ref}} \right)^m \exp \left[-\frac{Q_{mech}}{RT} \right]$$



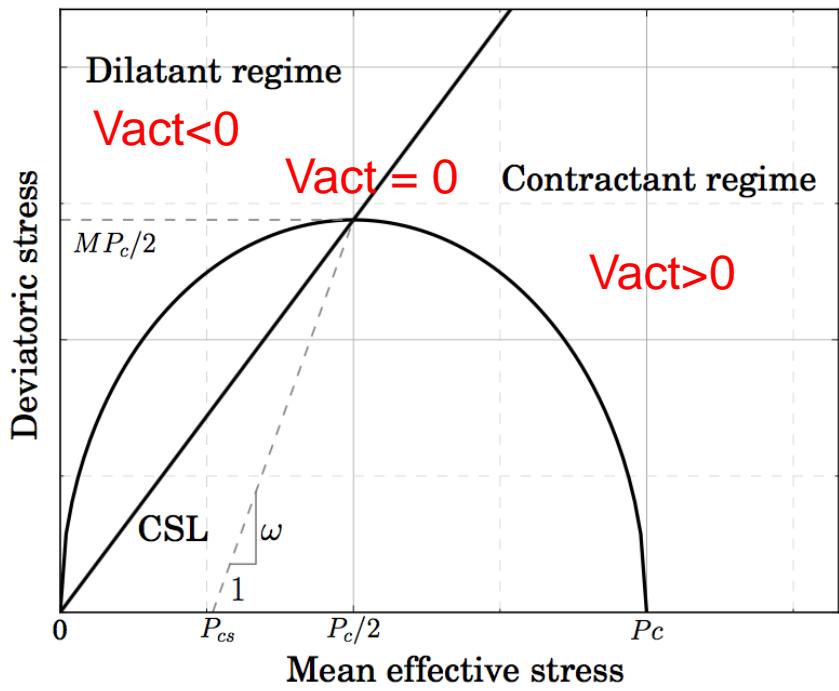
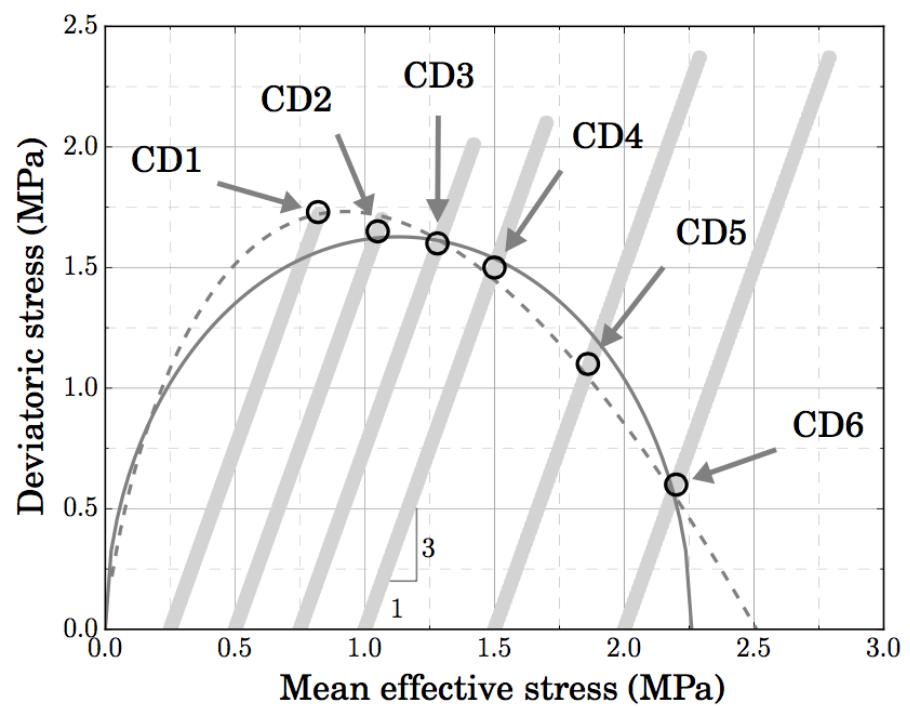
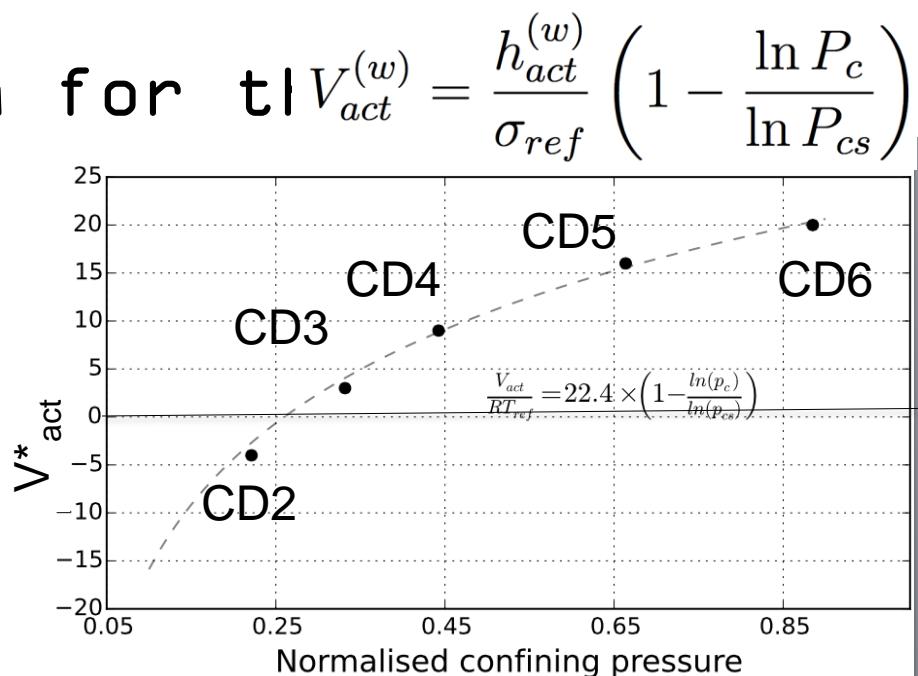
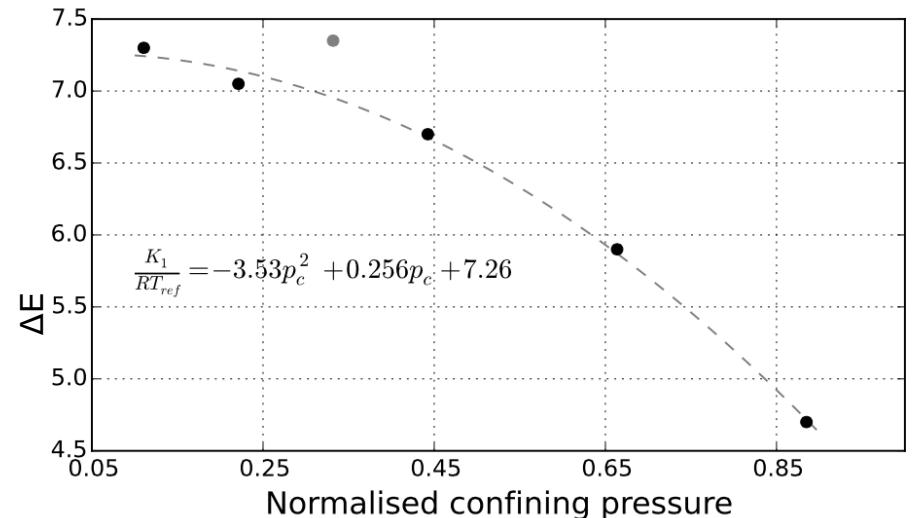
$$Q_{mech} = E + p_f V_{act}$$

$$E = E_0 + DE$$



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Retrieving information for tI



Modelling Physics + Chemistry of active faults

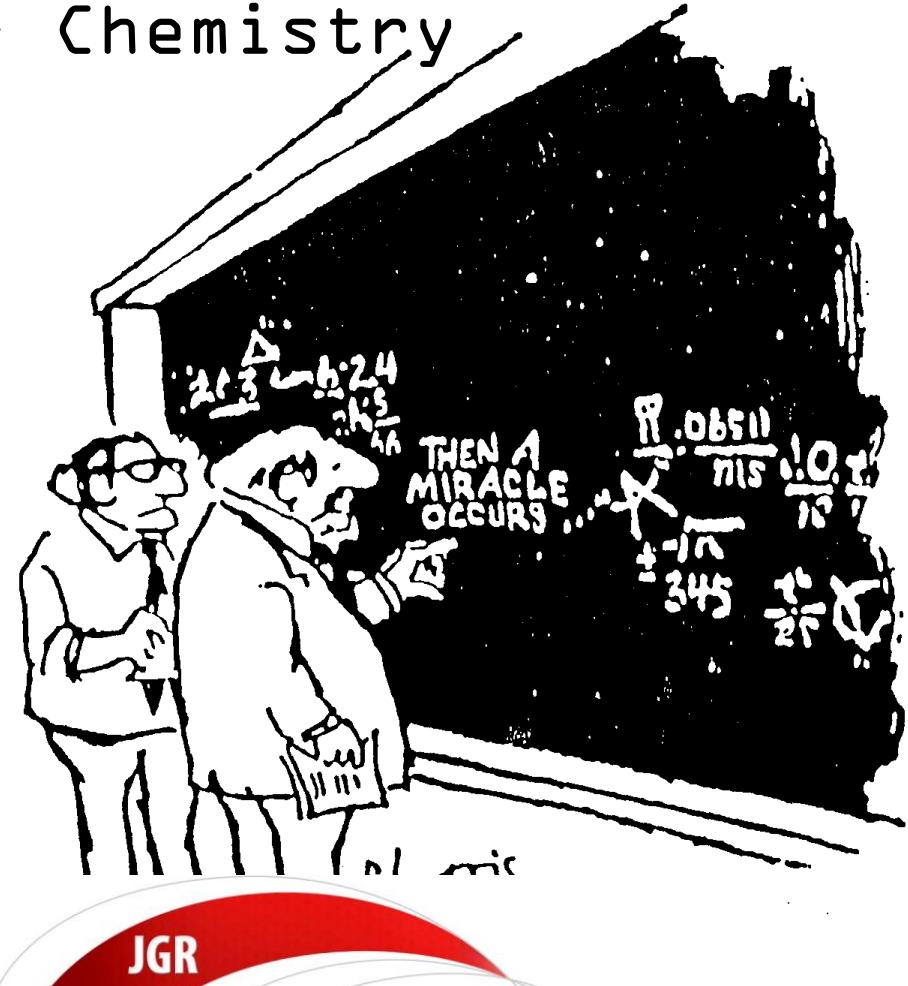
Part 1. Analytical Solution

- Multiple Steady States
- Asymptotic Solutions

Part 2. Transient Solutions

Spectral FE code SuCCoMBE

Part 3. Transition to Chaos



 AGU PUBLICATIONS

Journal of Geophysical Research: Solid Earth

RESEARCH ARTICLE

10.1002/2014JB011004

This is a companion paper
to Alevizos *et al.* [2014],
doi:10.1002/2013JB010070,
and Veveakis *et al.* [2014].

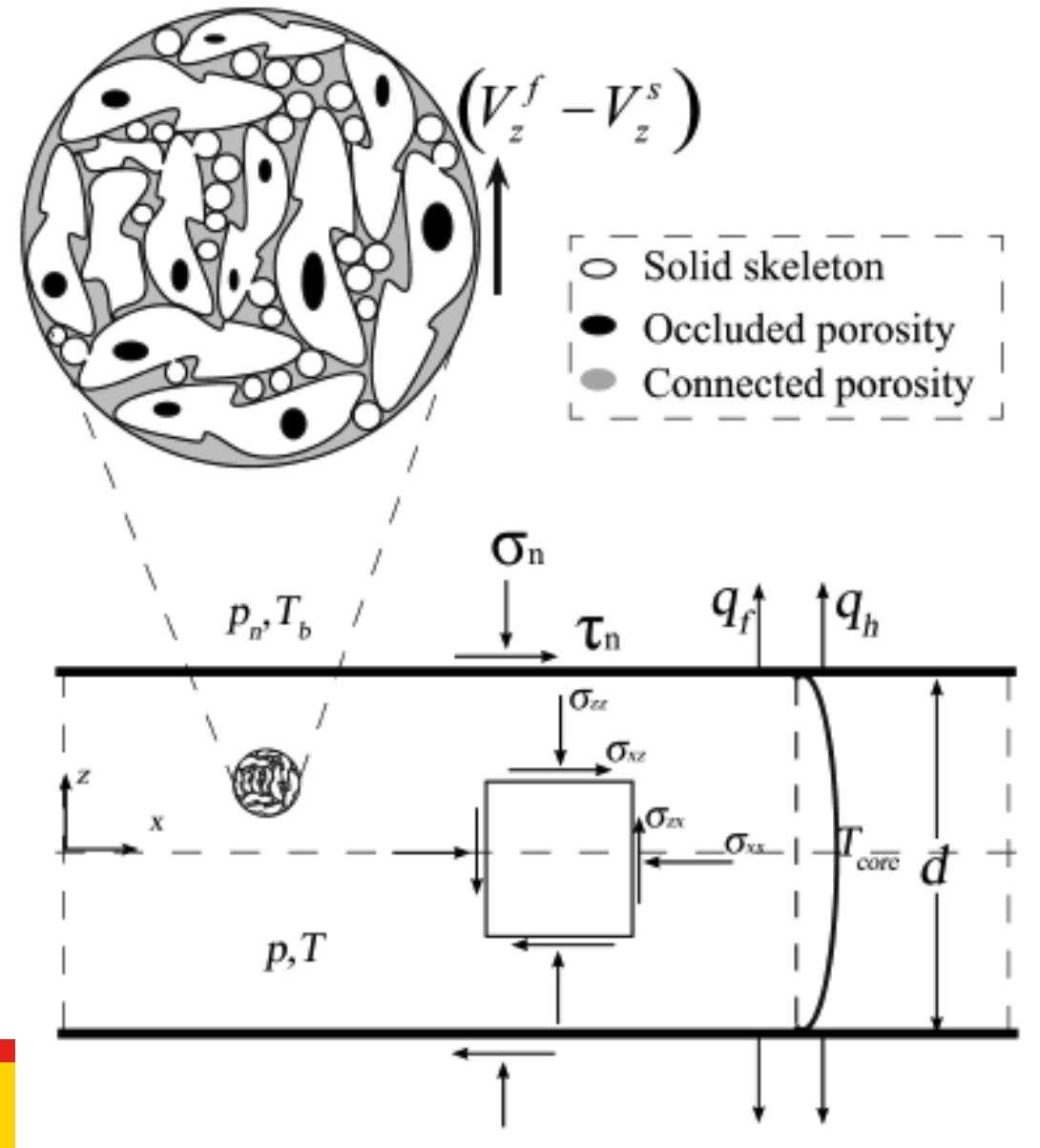
Thermo-poro-mechanics of chemically active creeping faults: 3. The role of serpentinite in episodic tremor and slip sequences, and transition to chaos

T. Poulet¹, E. Veveakis^{1,2}, K. Regenauer-Lieb^{1,3}, and D. A. Yuen^{4,5}



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Ingredients of the model



Basic physical mechanisms

- Creeping shear post-failure
- Shear Heating in creeping chemical process zone
- Chemical reaction/decomposition of the skeleton (excess pore pressure and chemical softening/hardening)
- Reverse chemical reaction



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The mathematical system

- Normalised and reduced system of equations

$$t = t_d(t), \quad S' = S'_n(t)$$

$$\frac{\partial DP}{\partial t} = \frac{\partial}{\partial z} \frac{1}{Le} \frac{\partial DP}{\partial z} + \frac{1}{m S'_n} \frac{\partial T}{\partial t} + (1 - f)(1 - s) m_r e^{\frac{Ar dT}{1+dT}}$$

$$\frac{\partial T}{\partial t} = \frac{\partial^2 T}{\partial z^2} + \frac{Gre}{\frac{DP Vact}{1+dT}} e^{\frac{aAr}{1+dT}} - (1 - f)(1 - s) e^{\frac{Ar dT}{1+dT}}$$

- Dimensionless Groups:

Lewis number

$$Le = \frac{\kappa_m \mu_f}{k_\pi \sigma'_n}, \quad \mu_r = \frac{(d/2)^2}{\kappa_m \sigma'_n} \frac{k_0}{\beta_f} e^{-Ar}, \quad Ar = \frac{E}{RT_c},$$

$$\delta = \frac{1}{m T_c}, \quad m = \frac{j k_m}{|\Delta H| (d/2)^2} \frac{e^{Ar}}{k_0 \rho_{AB}}, \quad Gr = \frac{\beta_T \tau_d \dot{\gamma}_0}{|\Delta H| k_0 \rho_{AB}}$$

Gruntfest number



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The mathematical system

- Normalised and reduced system of equations

$$t = t_d(t), \quad S' = S'_n(t)$$
$$\frac{\partial DP}{\partial t} = \frac{\partial}{\partial z} \frac{1}{Le} \frac{\partial DP}{\partial z} + \frac{1}{mS'_n} \frac{\partial T}{\partial t} + (1 - f)(1 - s) m_r e^{\frac{ArdT}{1+dt}}$$
$$\frac{\partial T}{\partial t} = \frac{\partial^2 T}{\partial z^2} + \frac{Gre}{1+dt} e^{-\frac{DP Vact}{1+dt}} e^{\frac{aAr}{1+dt}} - (1 - f)(1 - s) e^{\frac{ArdT}{1+dt}}$$

- Dimensionless Groups:

Lewis
number

$$Le = \frac{\text{heat diffusion}}{\text{mass diffusion}}$$

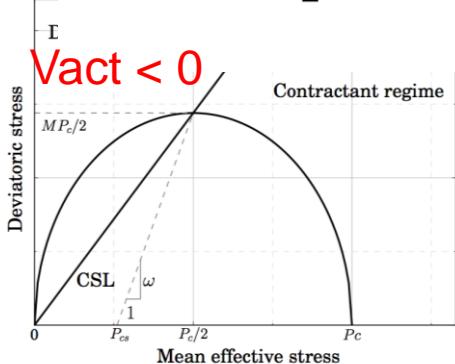
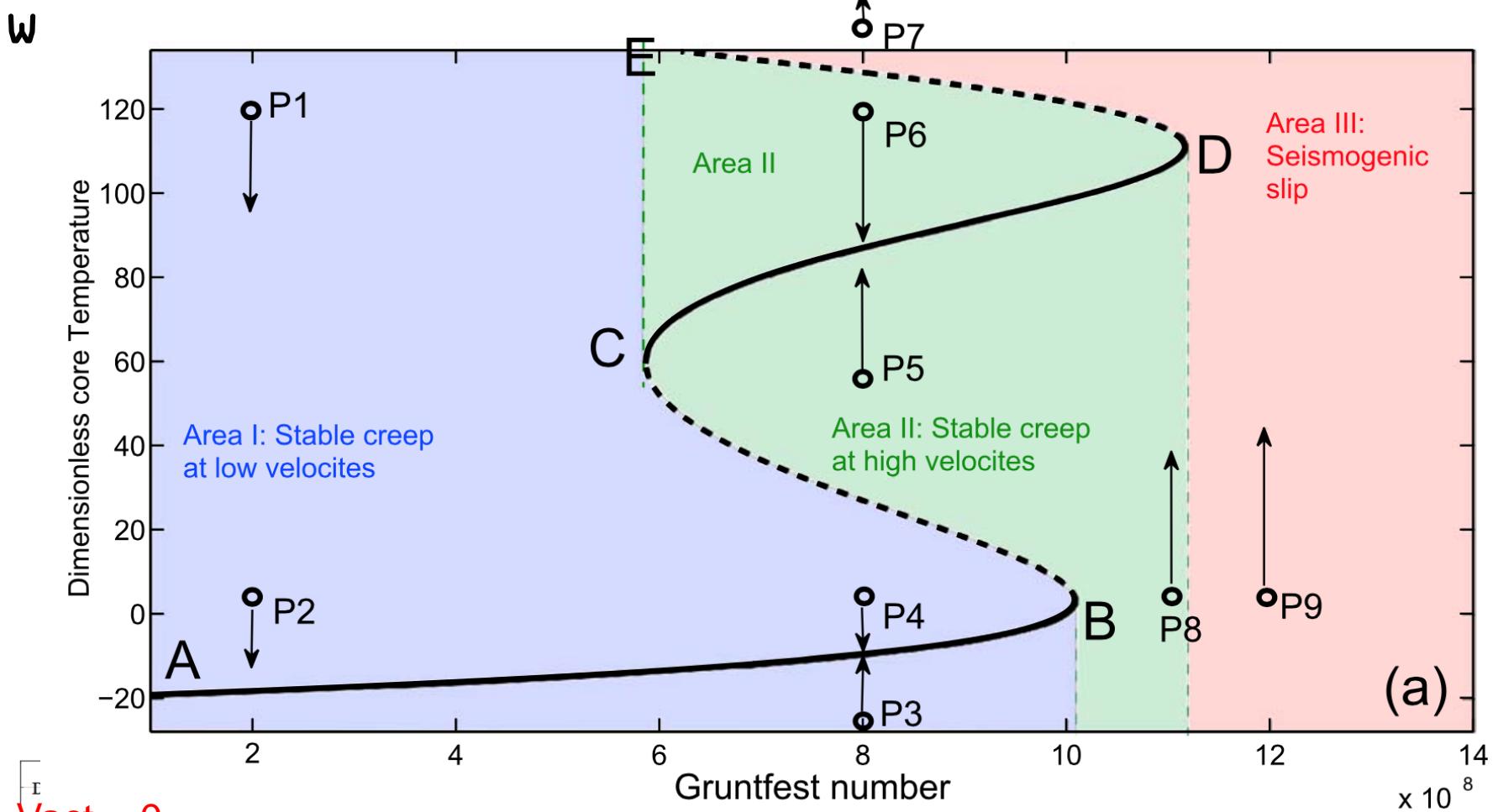
$$Gr = \frac{\text{char. time scale heat production}}{\text{char. time scale energy transfer}}$$

Gruntfest
number



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Vact < 0 : System's stability regimes

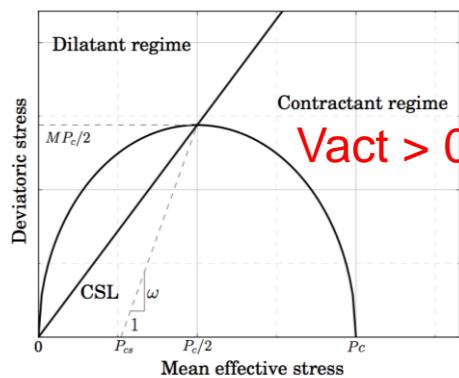
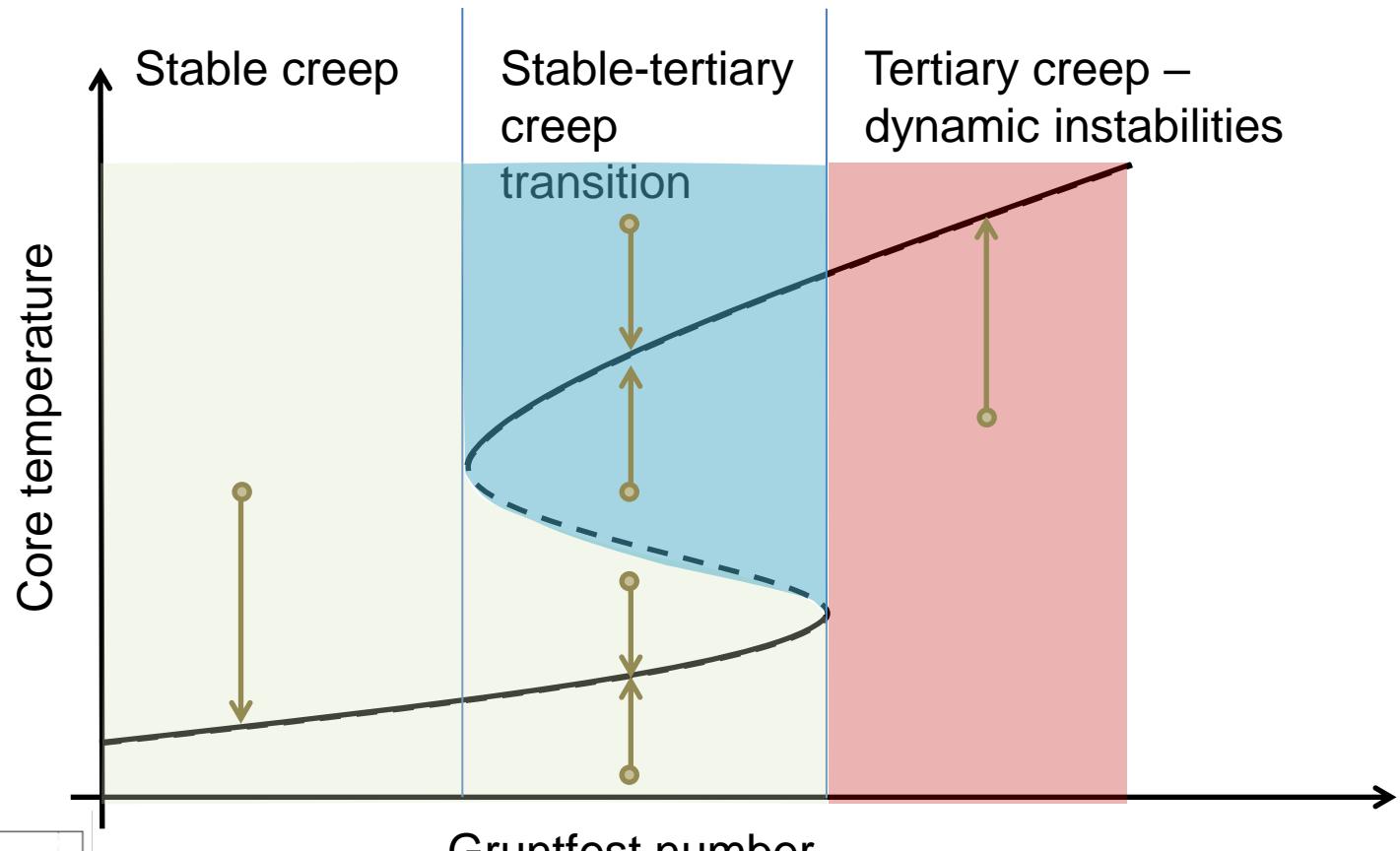


$$Gr = \frac{\text{char. time scale heat production}}{\text{char. time scale energy transfer}}$$



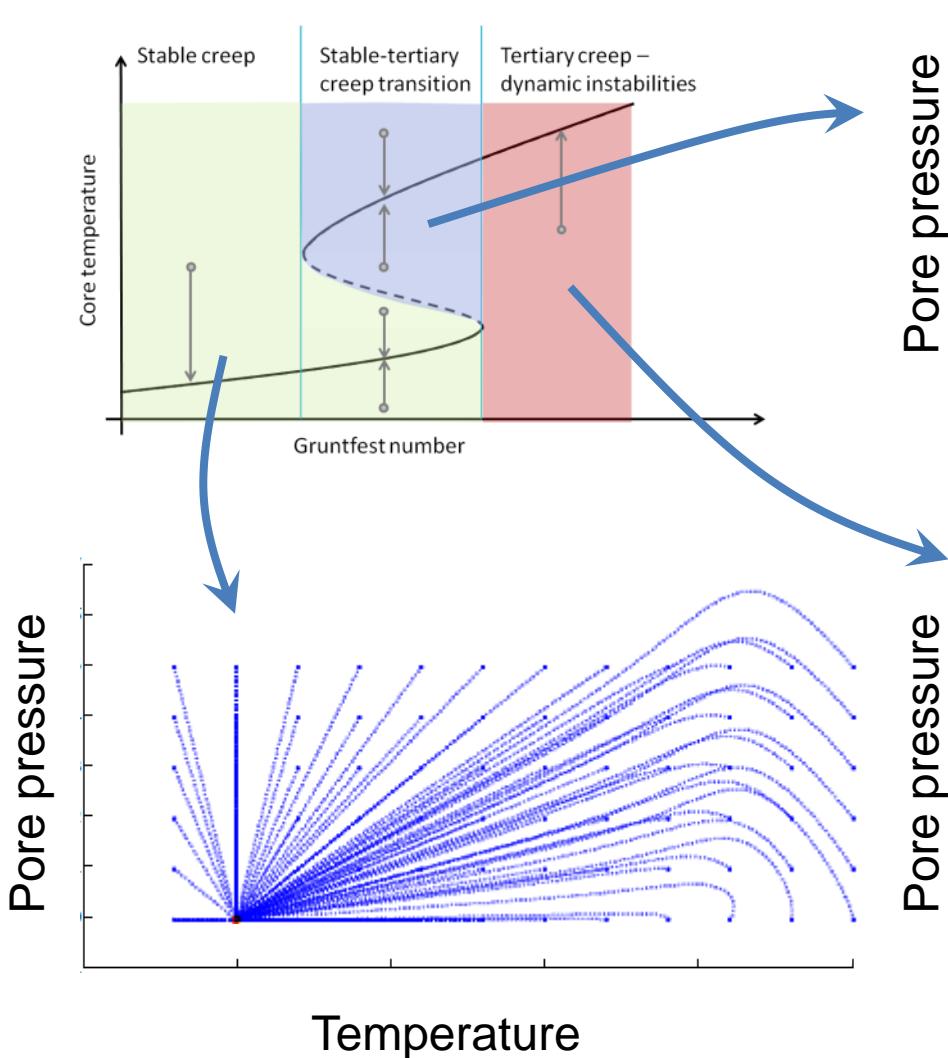
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$Vact > 0$: System's stability regimes w.r.t Gr

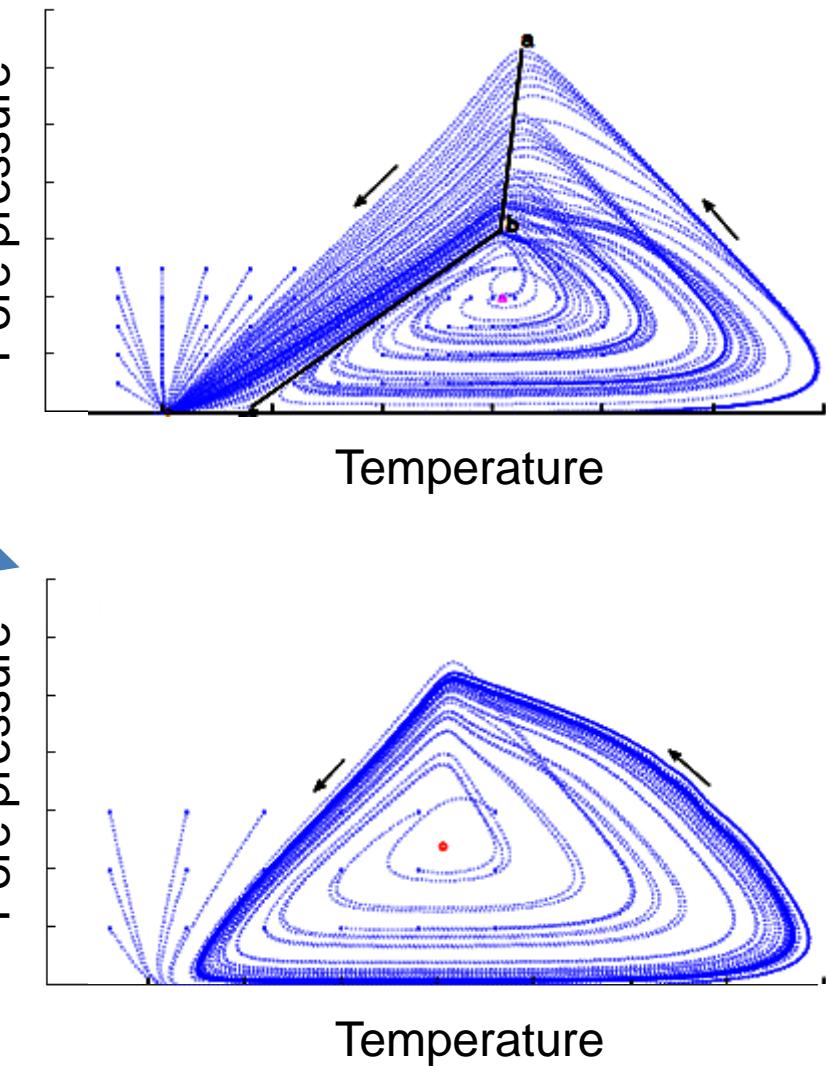


$$Gr = \frac{\text{char. time scale heat production}}{\text{char. time scale energy transfer}}$$

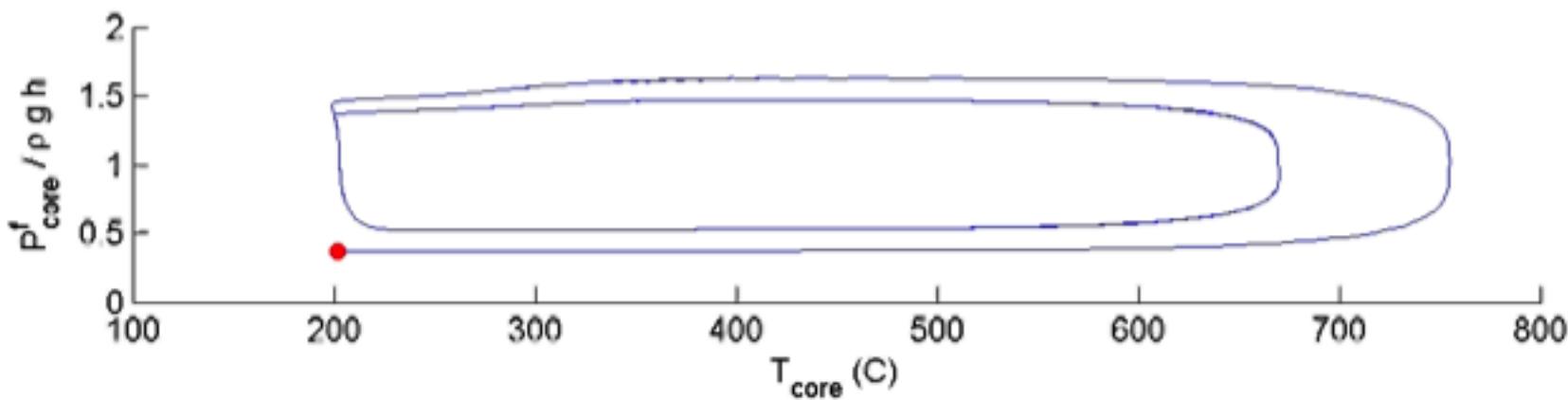
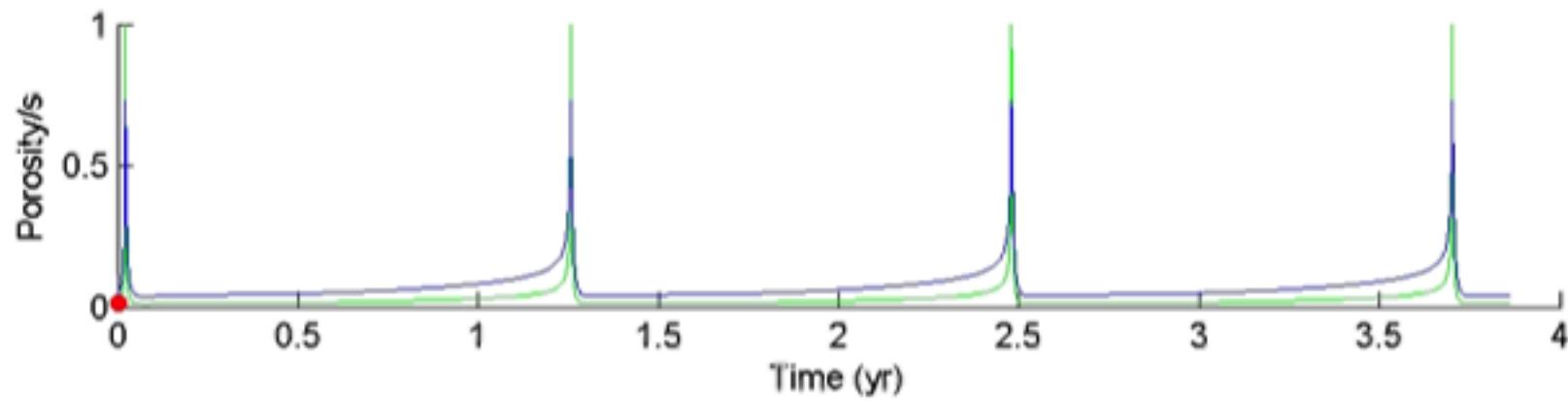
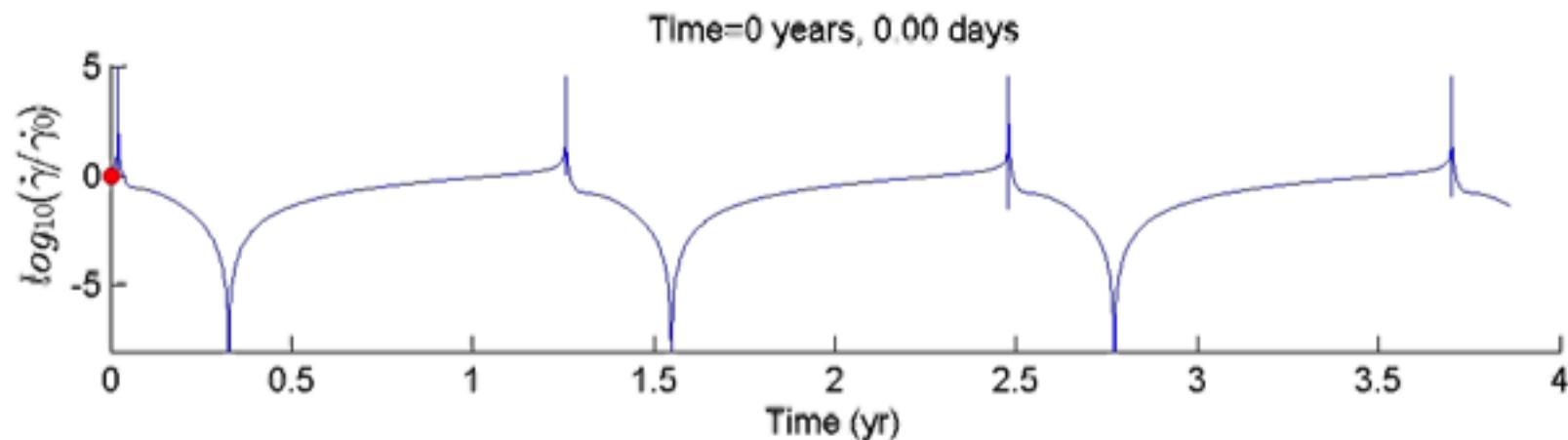
Phase diagrams



Natural localised instability

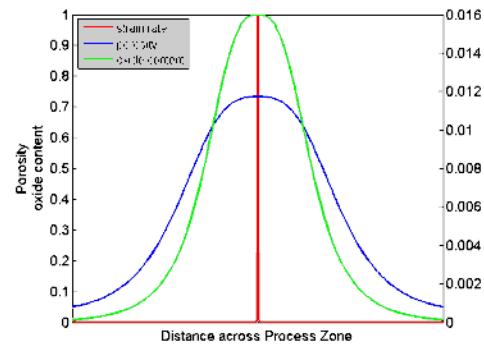
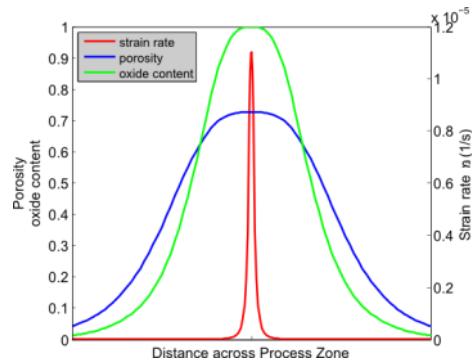


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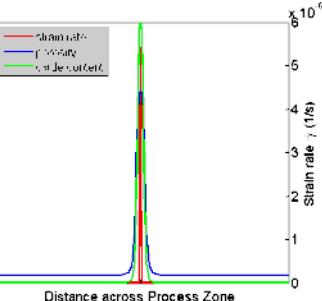
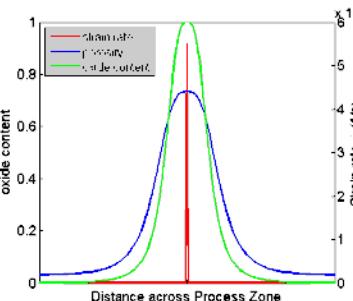
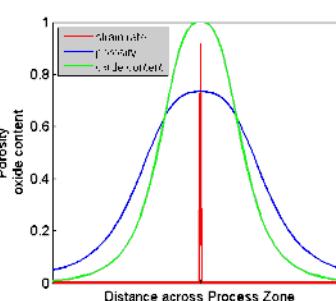
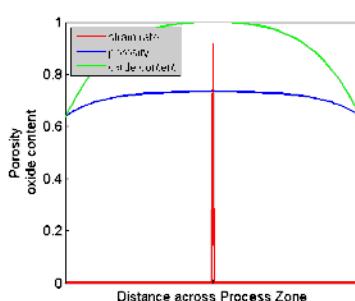
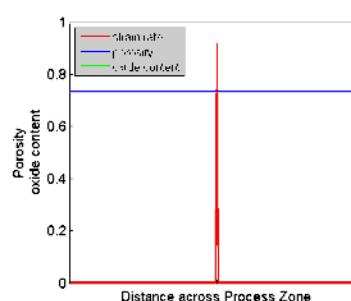


Spatial distributions

- Directly observable in the field
- Mechanical localisation: activation energy



- Chemical localisation: enthalpy of the reaction





Part III: Applications

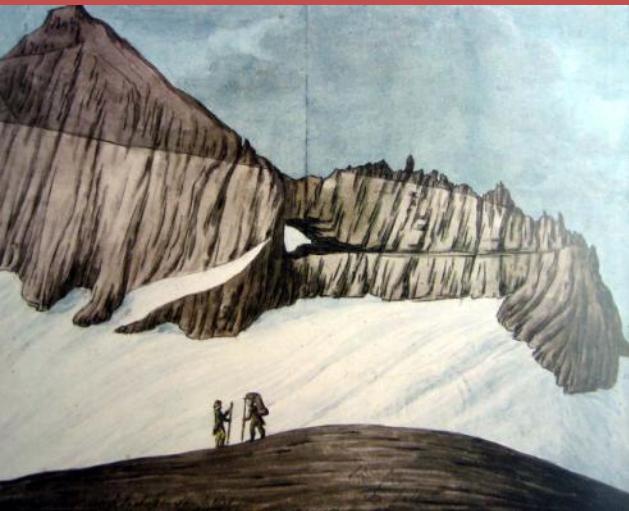
Fault Mechanics

Never Stand Still

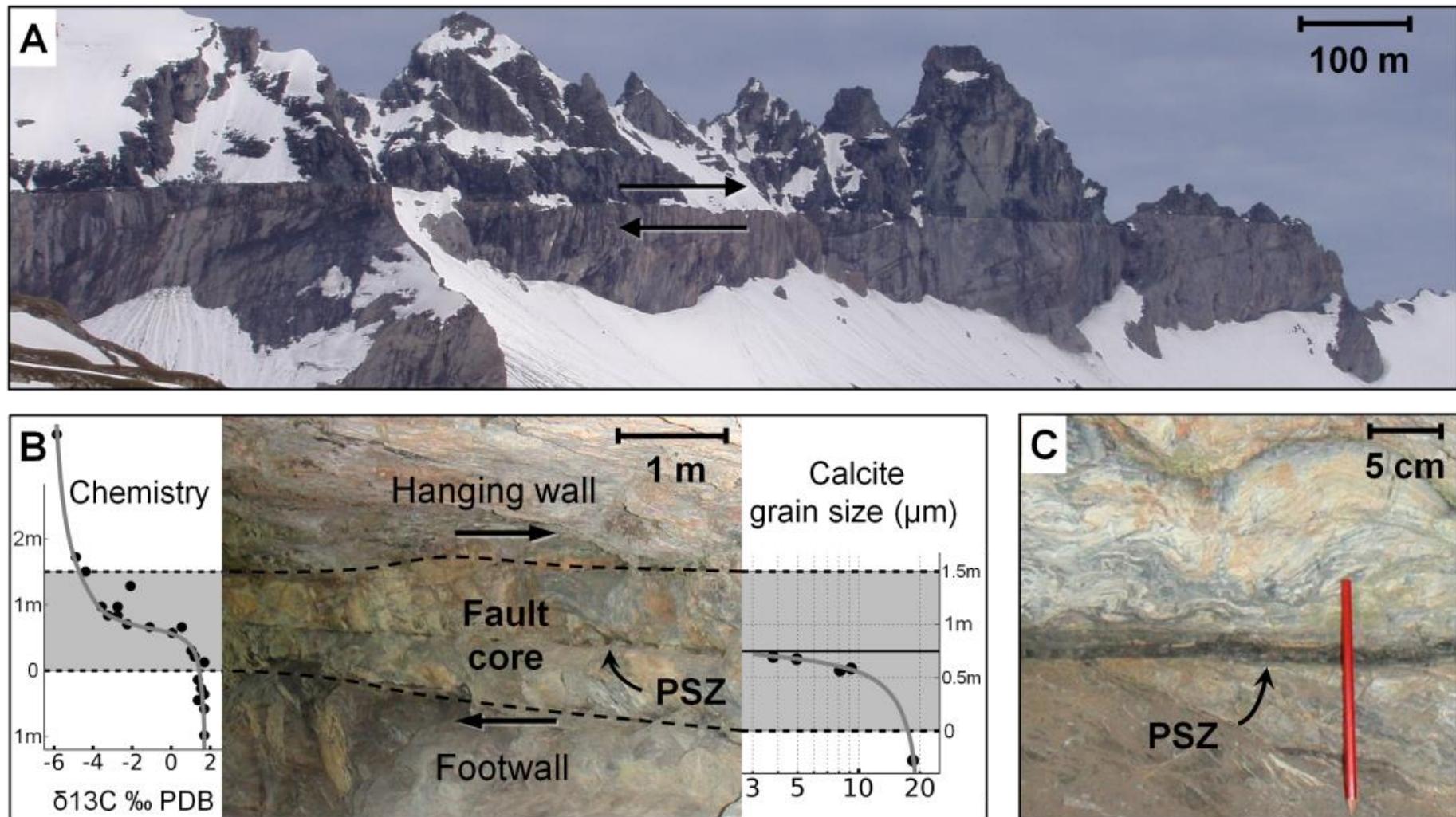
Engineering

The Glarus Thrust

A carbonate decomposition example

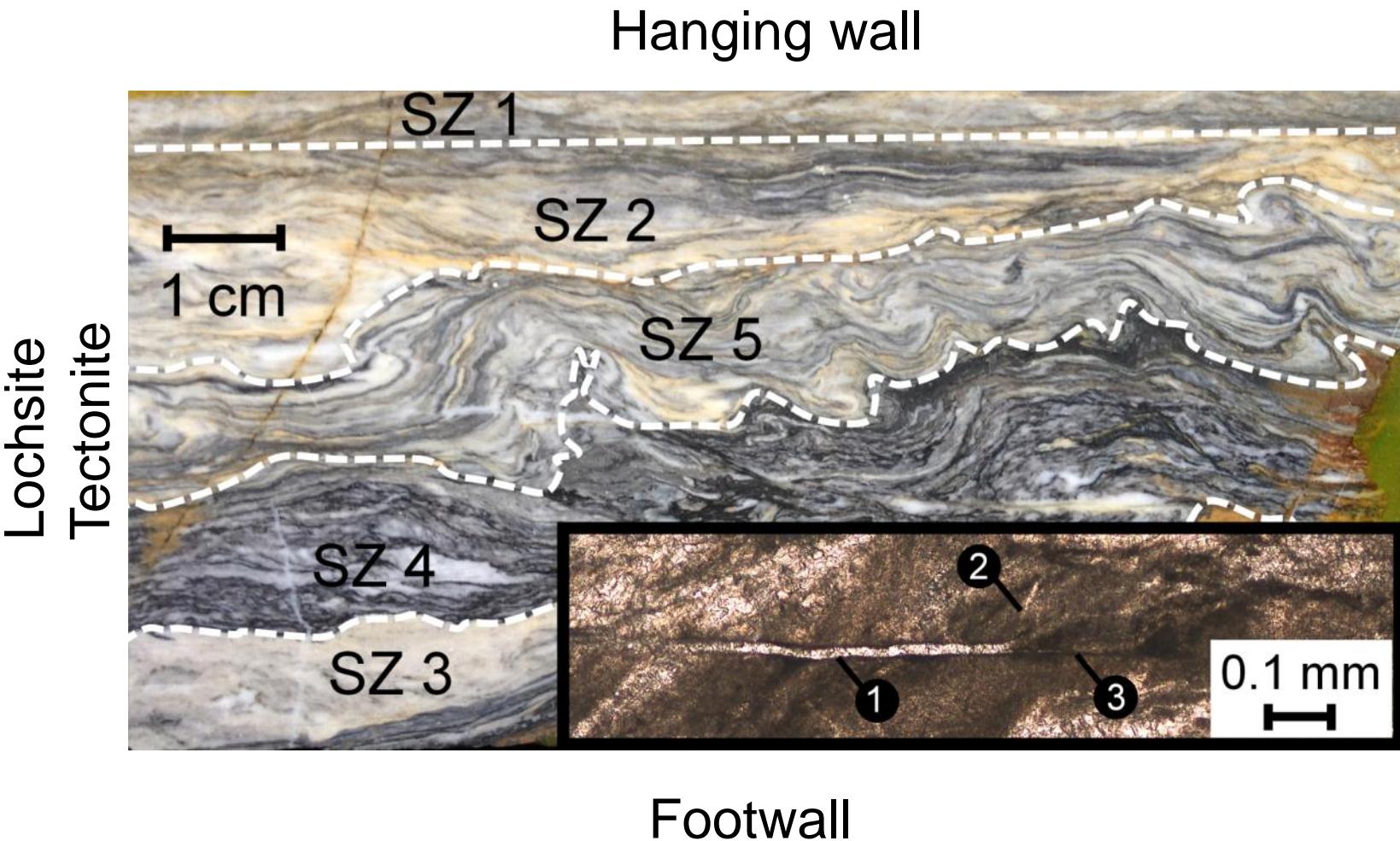


The Glarus Thrust, Switzerland



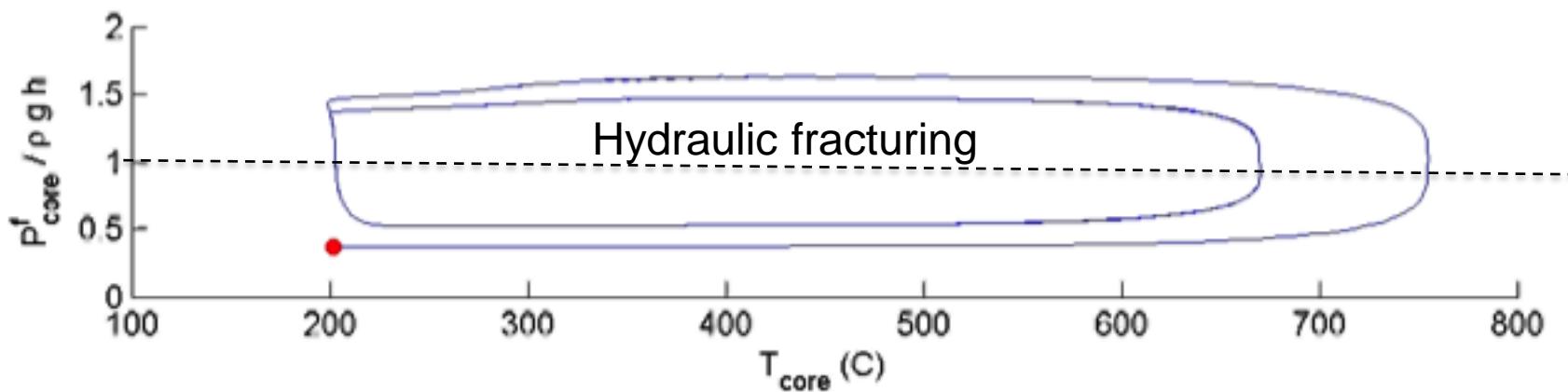
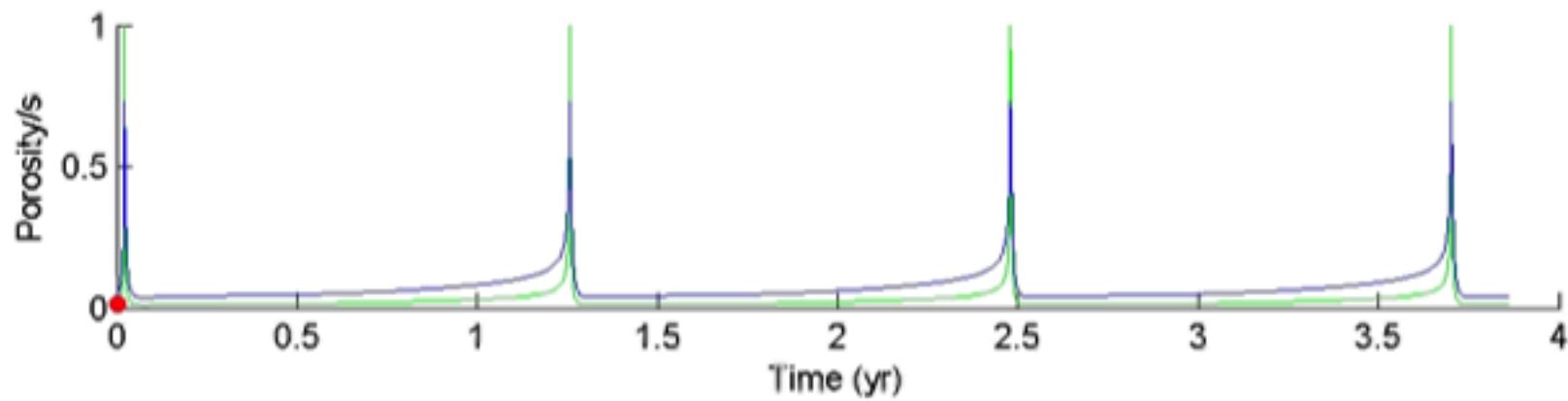
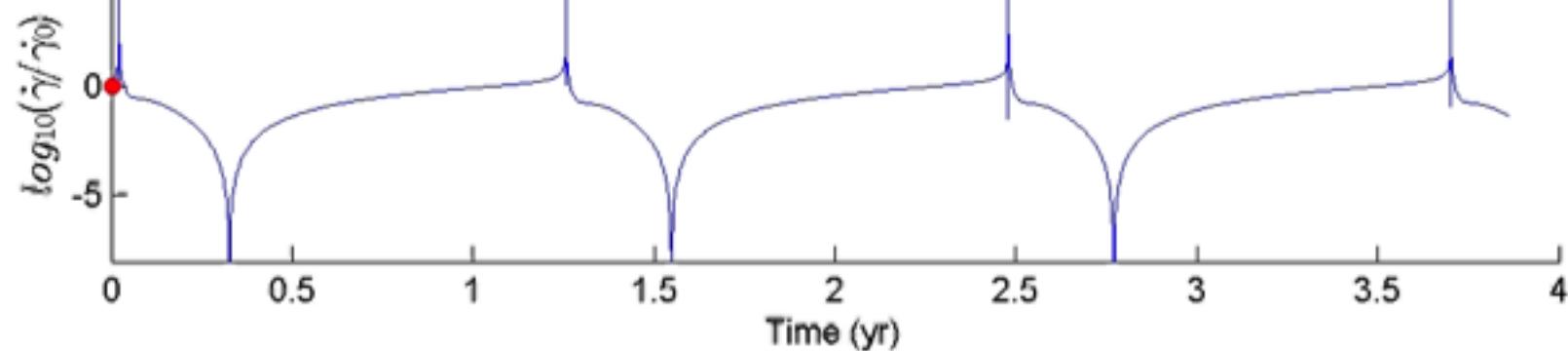
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The Glarus thrust - cyclicity

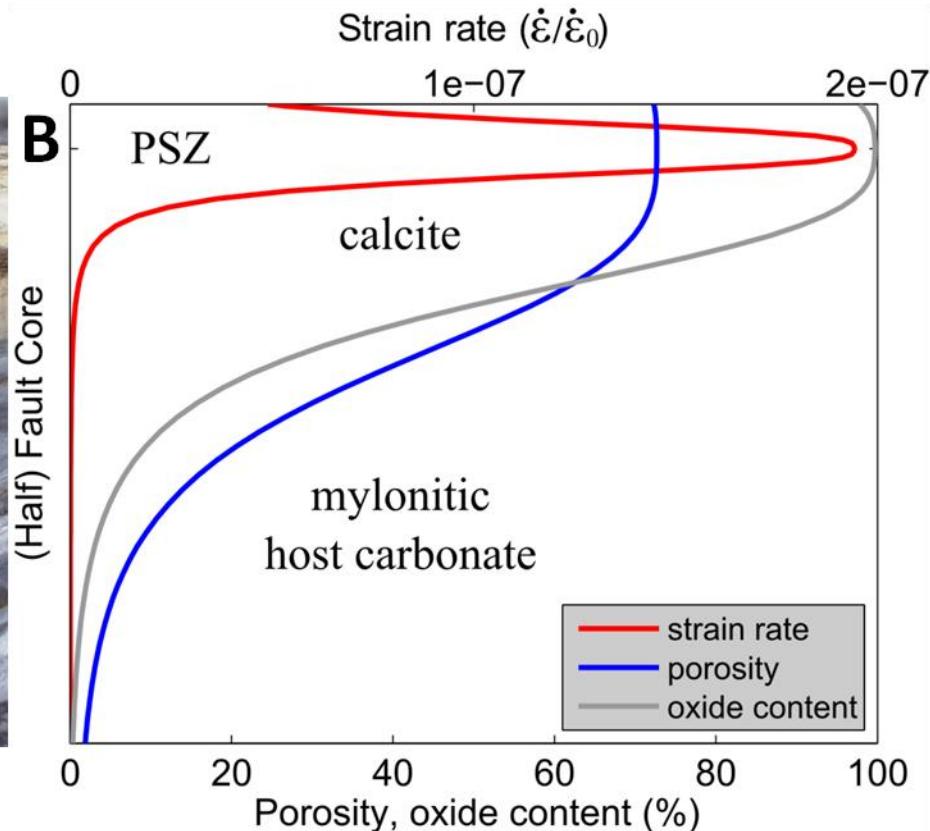
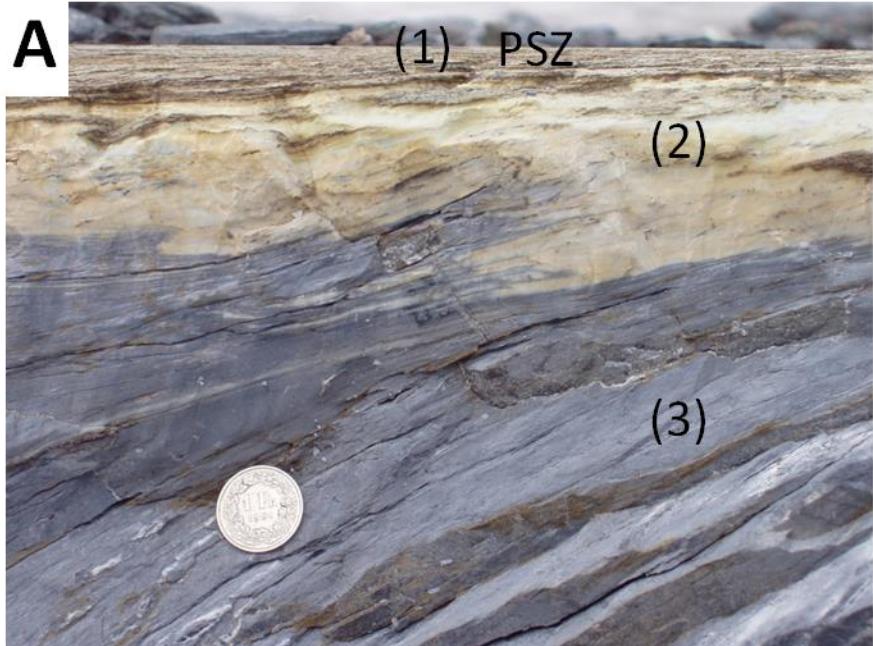


Temporal evolution

Time=0 years, 0.00 days



A Glarus pattern

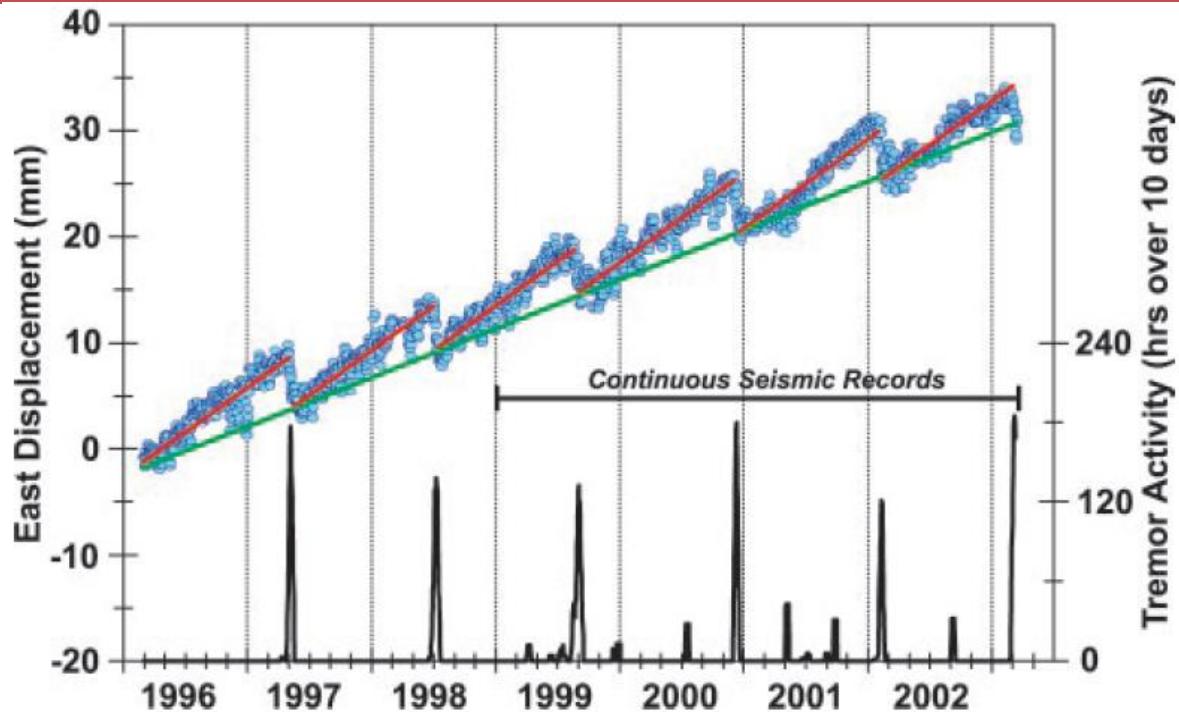
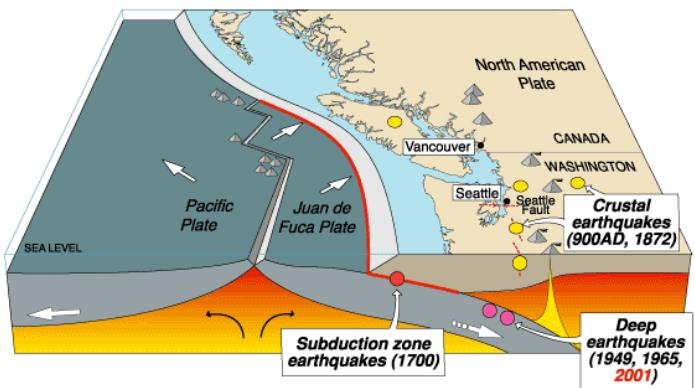


- Calcite as a cause (inverting activation enthalpy from thickness)
- Source of fluid explained (*in-situ* carbonate decomposition)

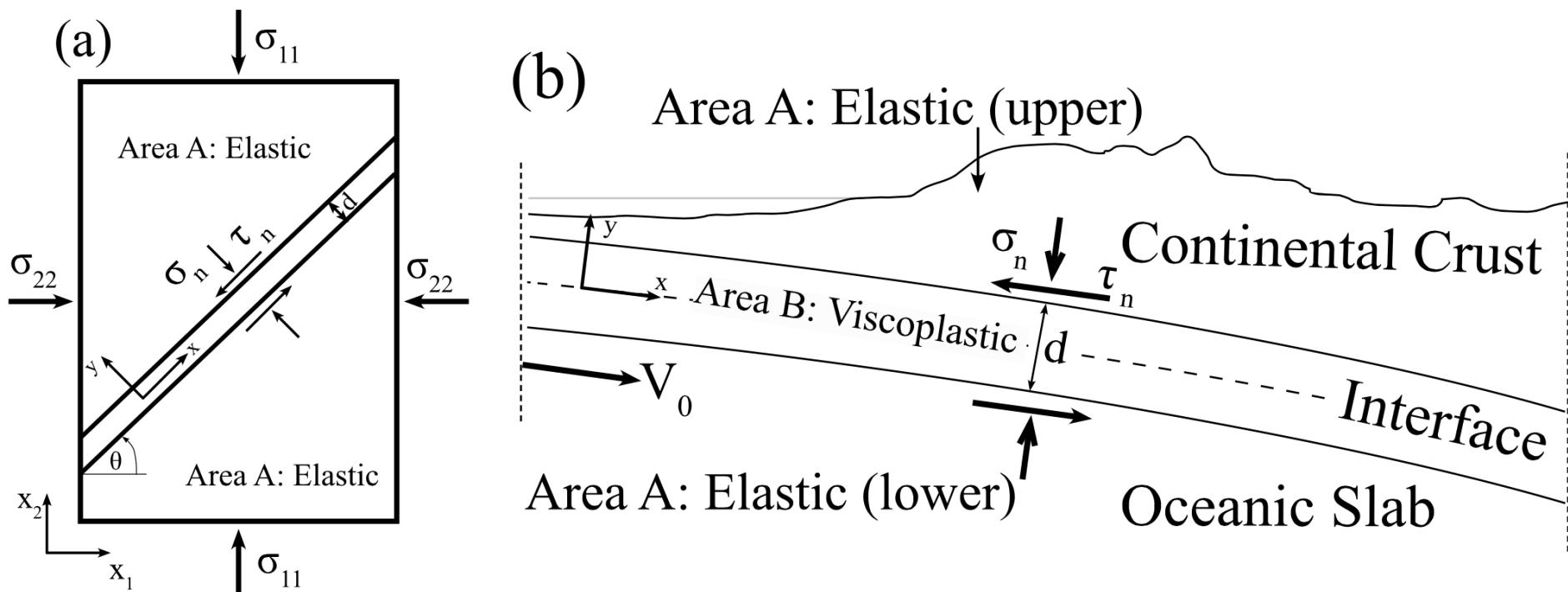
Subduction zones

Large scale modelling

Cascadia earthquake sources



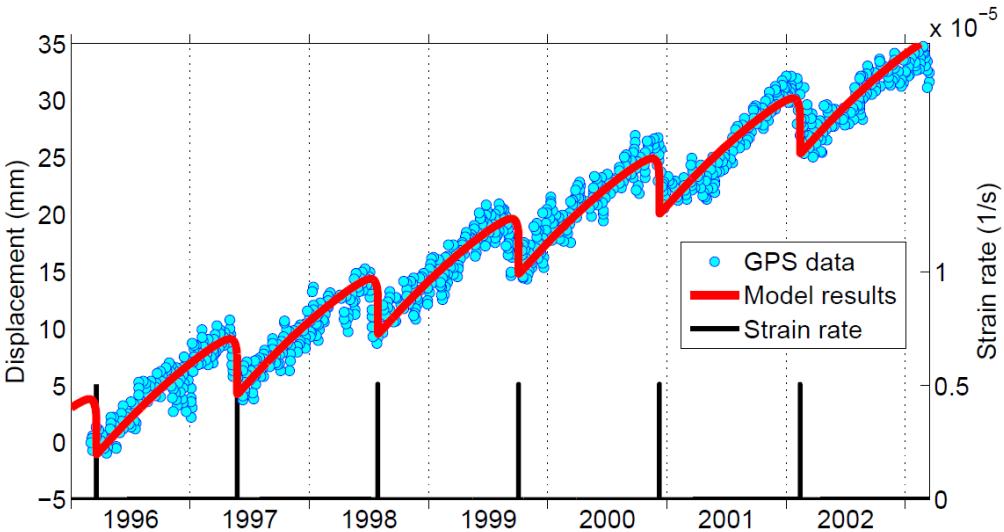
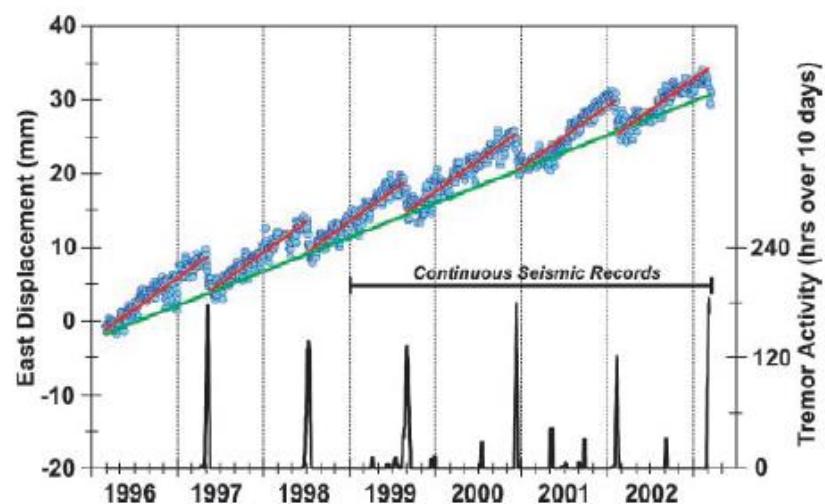
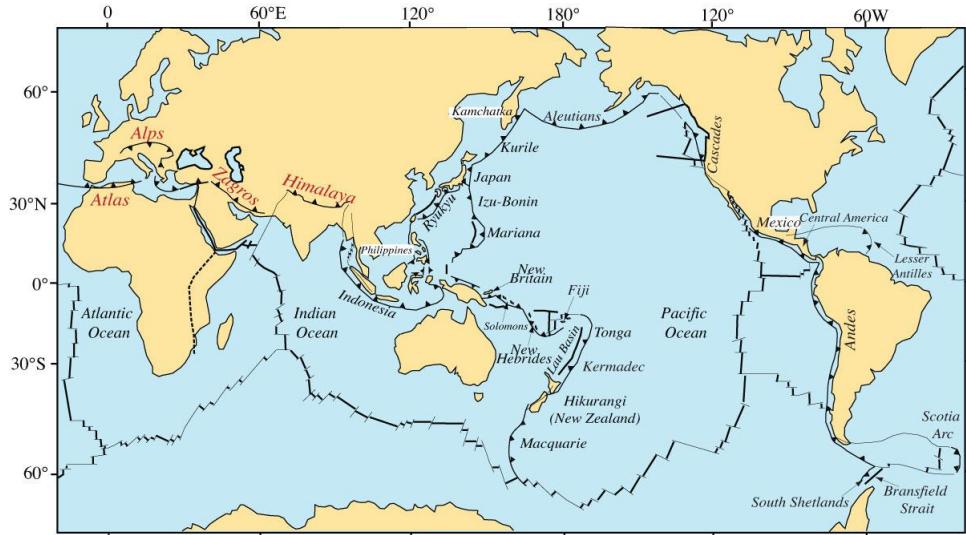
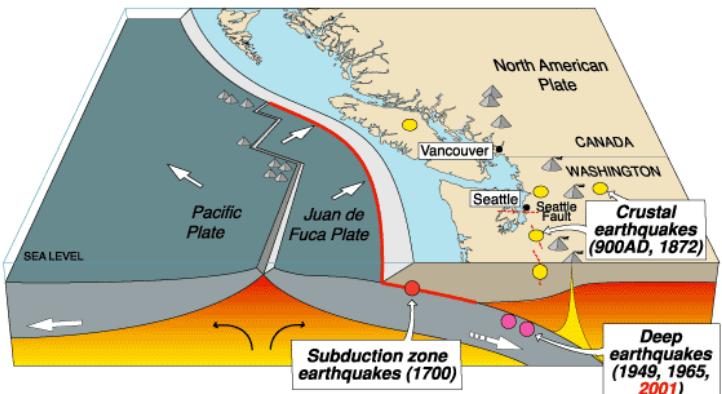
Modelling Subduction zones: Serpentinite dehydration oscillator



The problem: we require stress continuity at the shear band interface, because of the definition of the shear band as zone of velocity gradient discontinuities. So which BCs do we use?

Modelling subduction zones

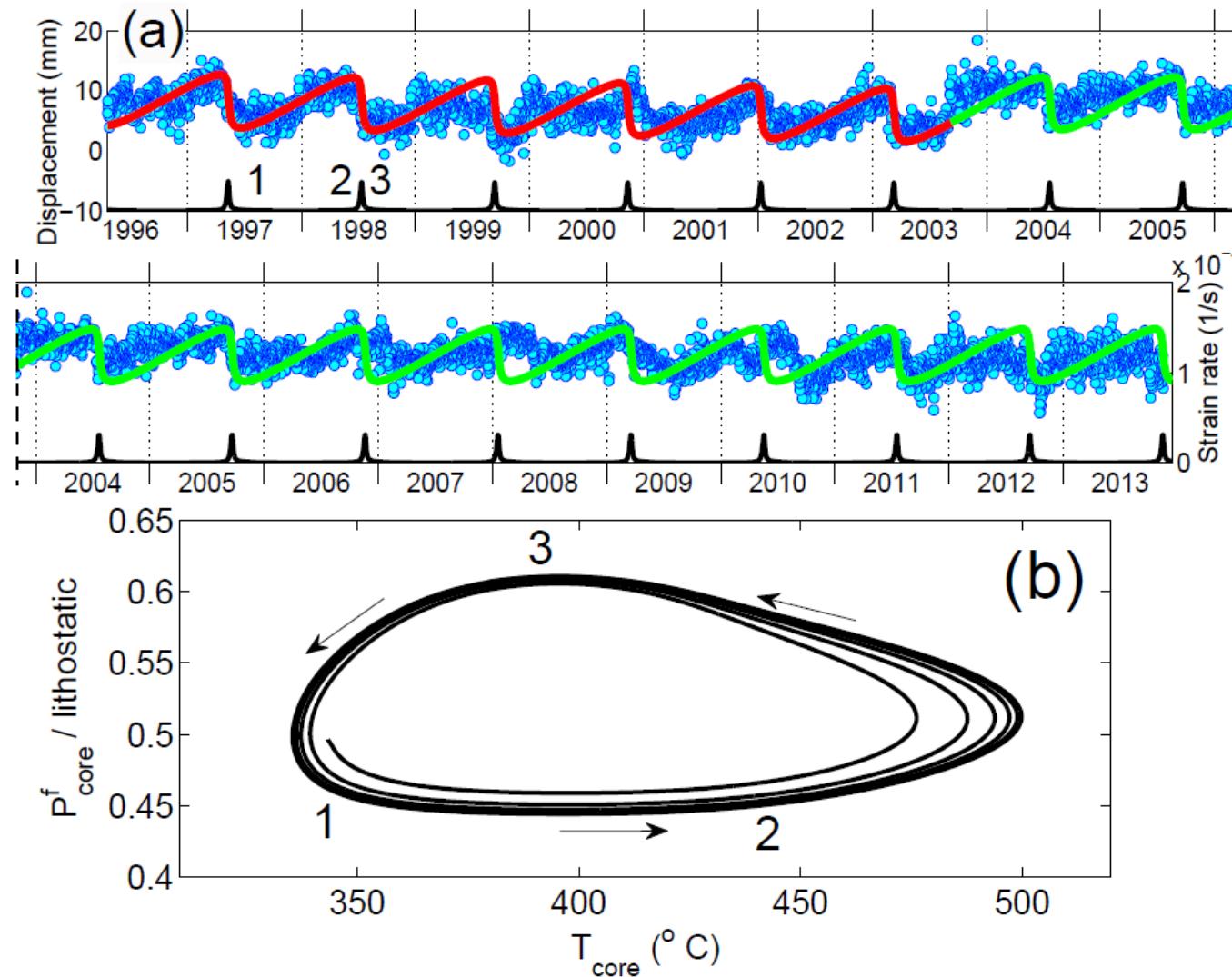
Cascadia earthquake sources



Rogers & Dragert, Science
(2003)

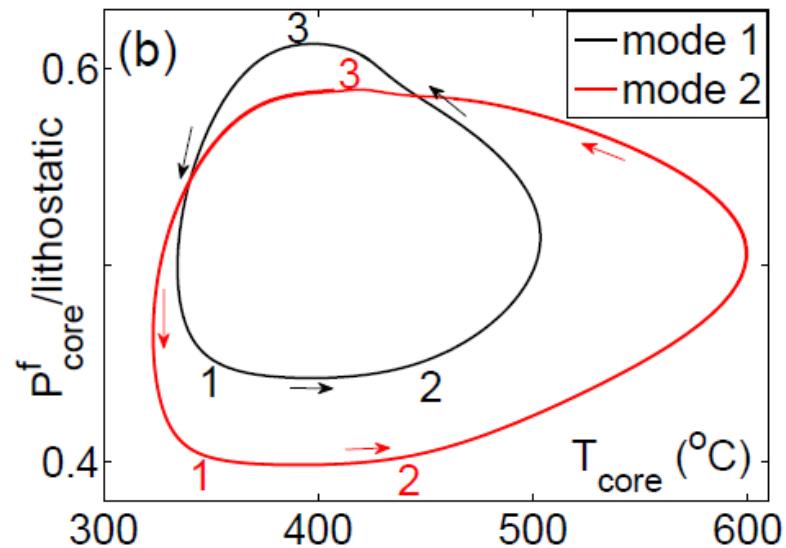
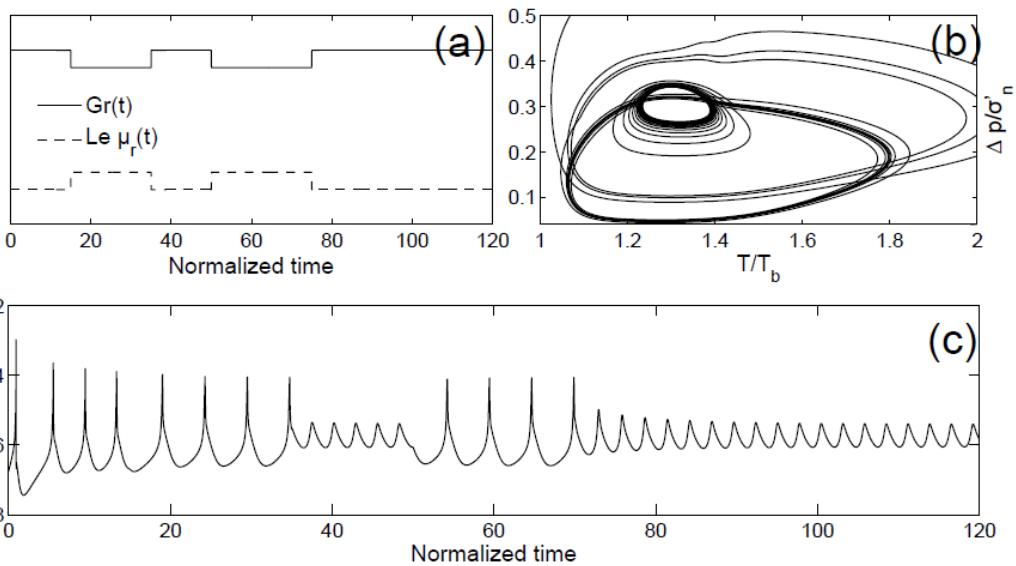
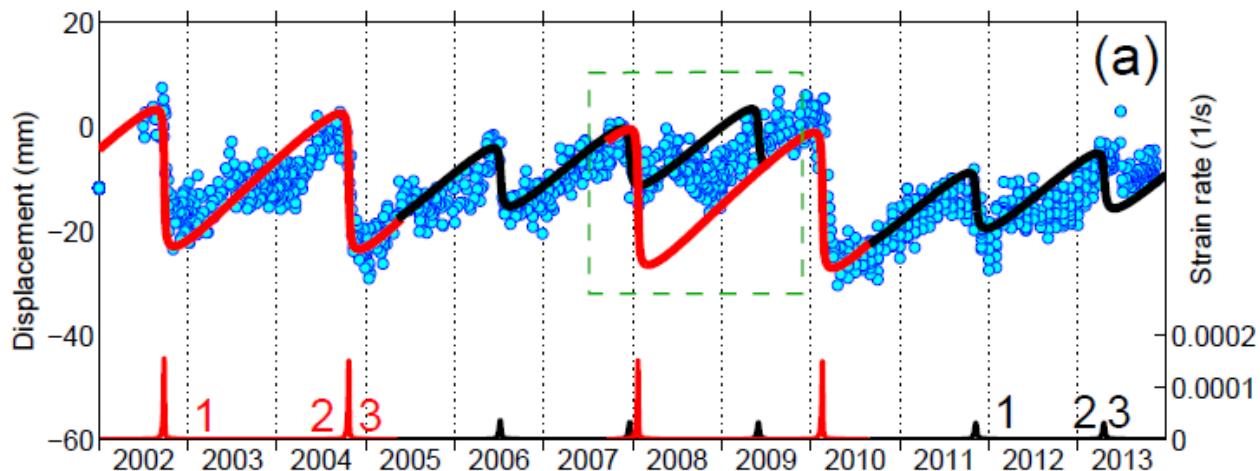
Alevizos et al, JGR
(2014)

Oscillator cycles, Earth's heartbeats

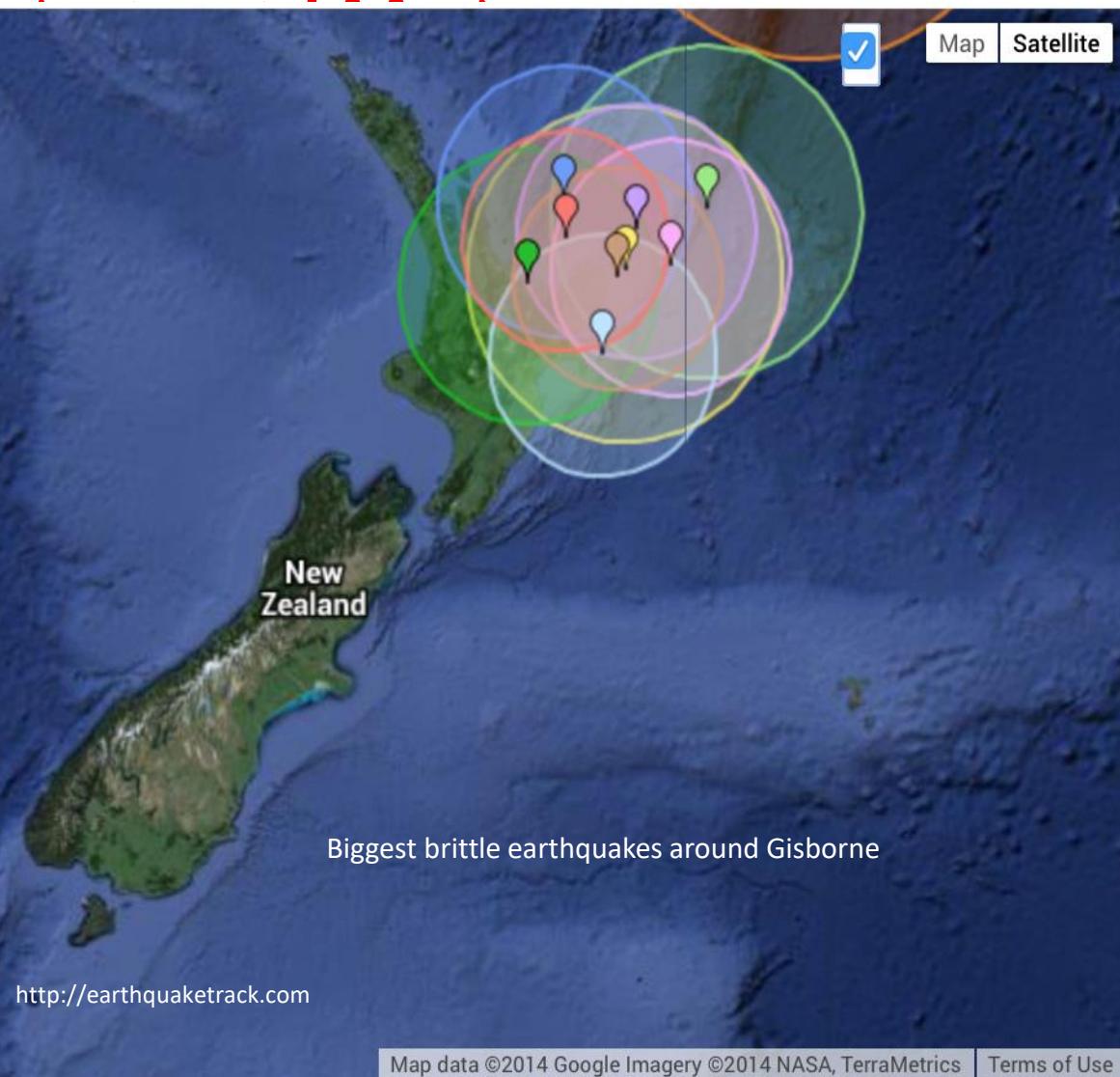


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Chaotic signals - Gisborne (New Zealand)



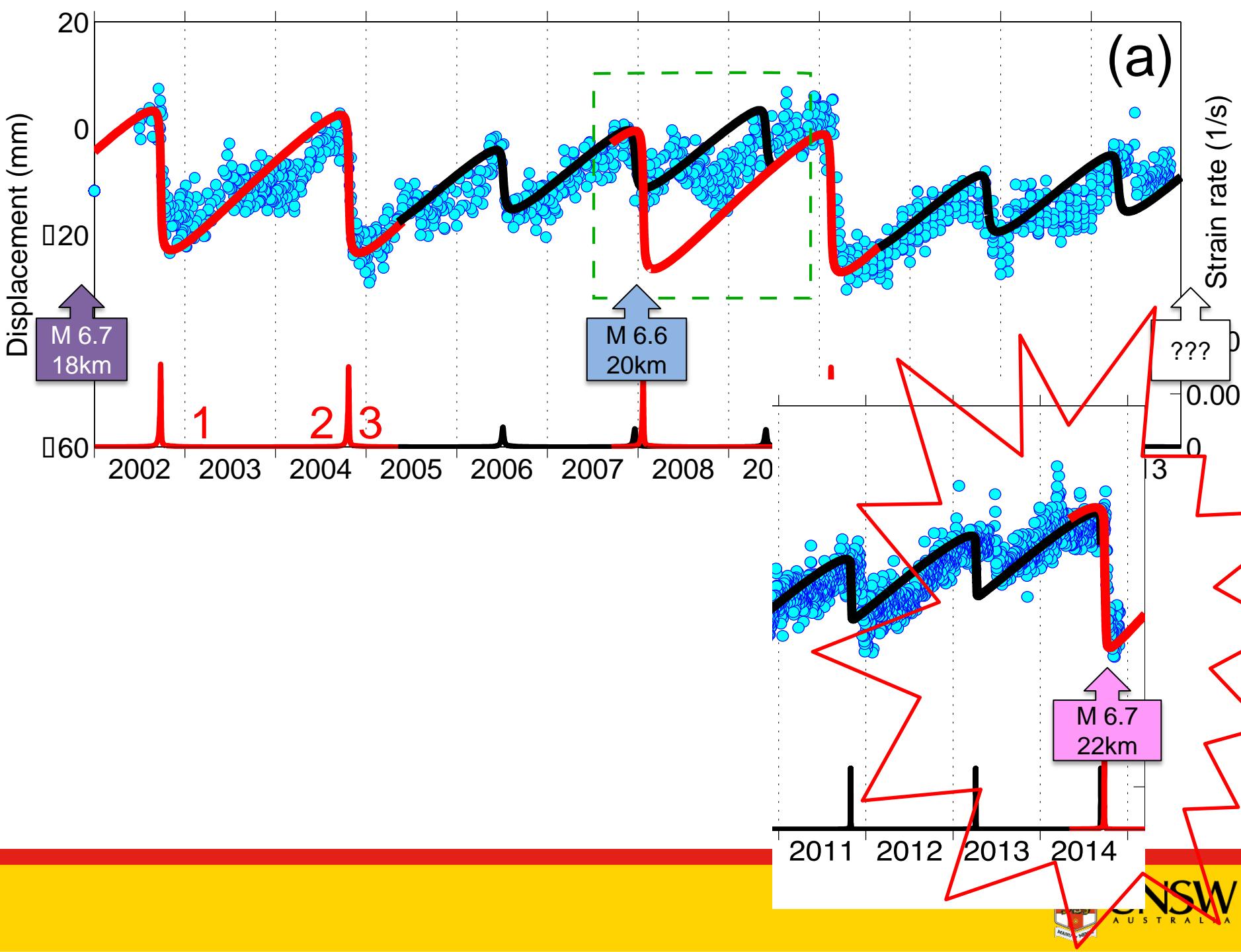
Can we predict brittle earthquakes? (NO)



- 37 years ago **7.7 magnitude**, 33 km depth
Whangarei, Auckland, New Zealand
- 20 years ago **7.1 magnitude**, 21 km depth
Gisborne, Gisborne, New Zealand
- 13 years ago **7.1 magnitude**, 33 km depth
Gisborne, Gisborne, New Zealand
- 28 years ago **6.8 magnitude**, 19 km depth
Whakatane, Auckland, New Zealand
- 30 years ago **6.8 magnitude**, 39 km depth
Whakatane, Auckland, New Zealand
- 13 years ago **6.7 magnitude**, 18 km depth
Gisborne, Gisborne, New Zealand
- 24 days ago **6.7 magnitude**, 22 km depth
Gisborne, Gisborne, New Zealand
- 7 years ago **6.6 magnitude**, 20 km depth
Gisborne, Gisborne, New Zealand
- 20 years ago **6.5 magnitude**, 28 km depth
Gisborne, Gisborne, New Zealand
- 20 years ago **6.5 magnitude**, 33 km depth
Whakatane, Auckland, New Zealand



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Matching observations

- Intern. Ocean Discovery Program
- Japan Trench Fast Drilling Project



Nature news, Dec 2013:

“The localization of deformation onto a limited thickness (~5 meters) of pelagic clay is the defining characteristic of the shallow earthquake fault” (Chester et al / Science 2013). “That’s just weird” says Emily Brodsky (UC Santa Cruz)

AGU PUBLICATIONS

Journal of Geophysical Research: Solid Earth

RESEARCH ARTICLE

10.1002/2014JB011004

This is a companion paper
to Aleivizos et al. (2014),
doi:10.1002/2013JB010070,
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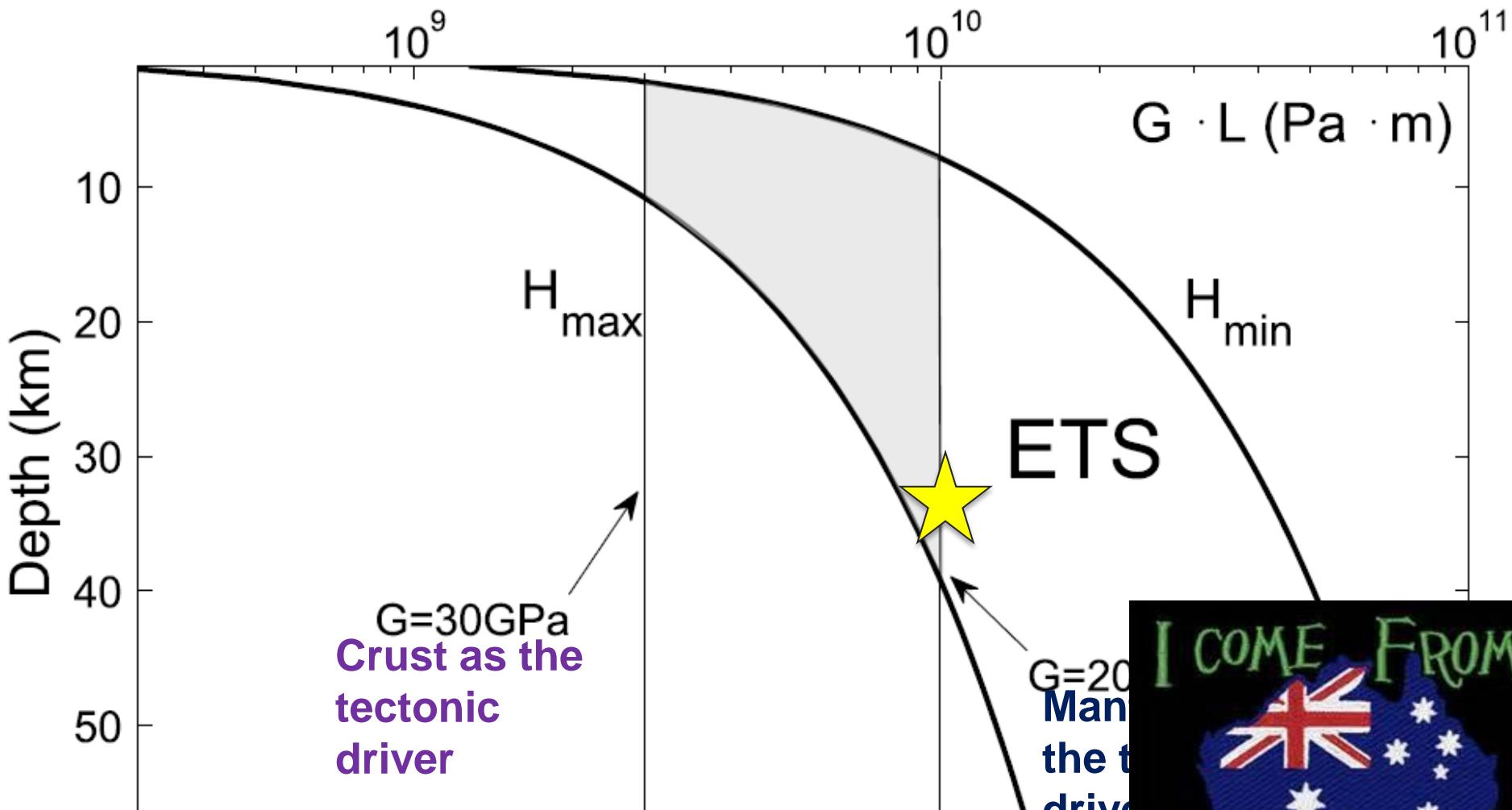
Thermo-poro-mechanics of chemically active creeping faults:
3. The role of serpentinite in episodic tremor and slip
sequences, and transition to chaos

T. Poulet¹, E. Veveakis^{1,2}, K. Regenauer-Lieb^{1,3}, and D. A. Yuen^{4,5}

Table 3. Material Parameters Inverted From the ETS Sequences, After Fitting the GPS Data^a

Parameter	Units	ALBH	GISB 1	GISB 2
$\dot{\gamma}_0$	s^{-1}	200	230	230
d	m	6.4	6.4	6.4
$\bar{\sigma}'_n$	MPa	49	49	74
$\beta_T \bar{\tau}_n$	MPa	0.3	0.26	0.20
k_F	s^{-1}	10^8	10^8	10^8
Q_F	kJ/mol	114	114	114
k_R	s^{-1}	10^{-2}	10^{-2}	10^{-2}
ΔH	kJ/mol	80	80	80
Q_R	kJ/mol	34	34	34

ETS location and tectonic driver: Cascadia



TECTONICS DOWN-UNDER!



I COME FROM
THE LAND
DOWN UNDER

More is less



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<http://github.com/pou036/redback.git>

Some of our work on the topic

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- Poulet T., M. Paesold and E. Veveakis, 2016. Multiphysics modelling for fault mechanics using REDBACK: A parallel open-source simulator for tightly coupled problems, *Rock Mechanics and Rock Engineering*, *in press (available online)*, doi: 10.1007/s00603-016-0927-y
- Peters M., M. Herwegh, M. K. Paesold, T. Poulet, K. Regenauer-Lieb and E. Veveakis, 2016. Boudinage and folding as the same energy attractor of ductile deformation, *J. Geophys. Res.*, 121, 3996–4013, doi: 10.1002/2016JB012801
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